

Tectono-Magmatic Evolution of Atlantic Type Rifted Margins

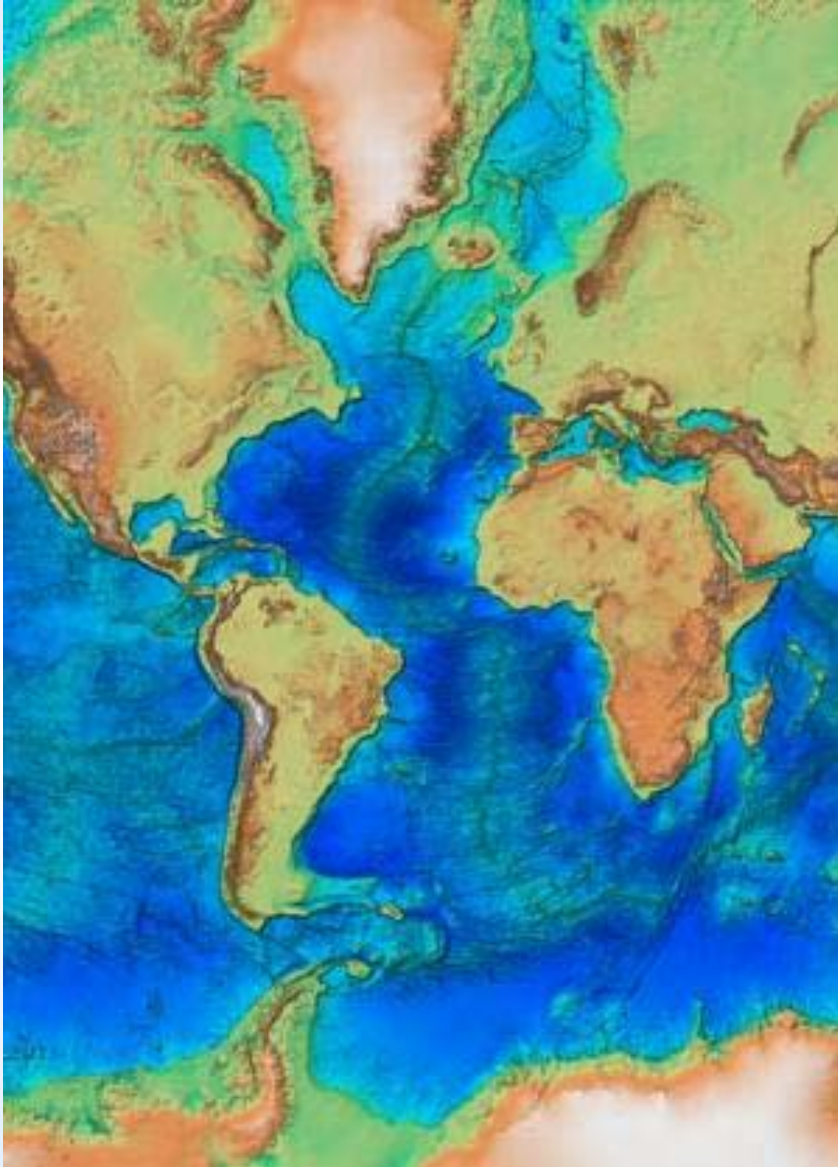
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Magmatic Rifting & Active Volcanism Conference
10-13/1/2012 Addis Ababa



Image: GeopoliticusChild by Salvador DALI



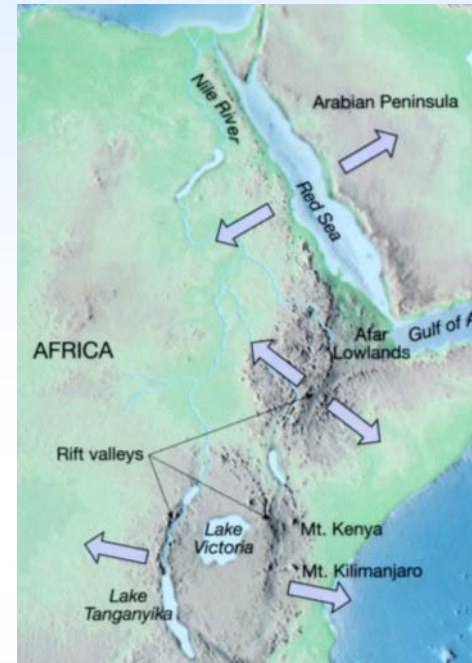
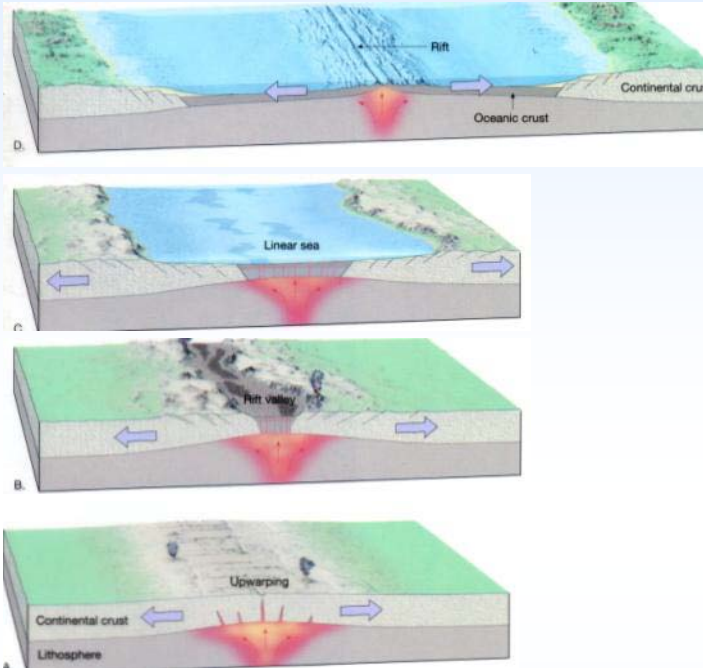
Major questions addressed in this presentation:

- ***What makes plates so weak that they can break?***
- ***Where, when and how does continental breakup occur?***
- ***What are the processes controlling rifting, continental breakup and sea floor spreading?***

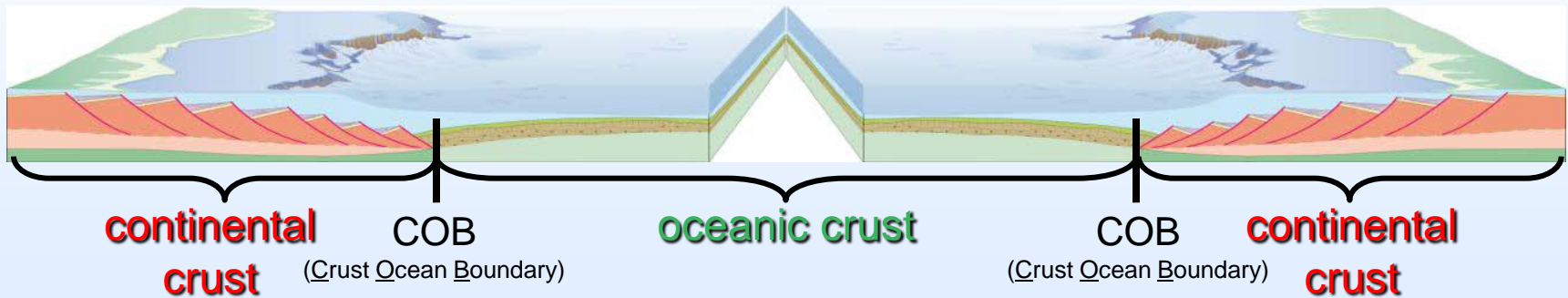
The “text book” rifted margin

A model inspired from the E-African and Red Sea rift system

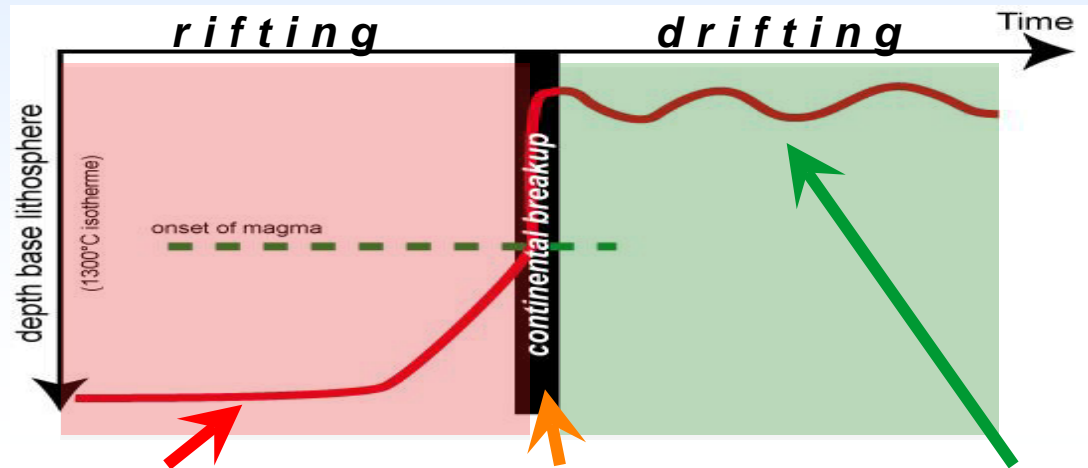
evolution of a rift system



Architecture of the “text book” rifted margin



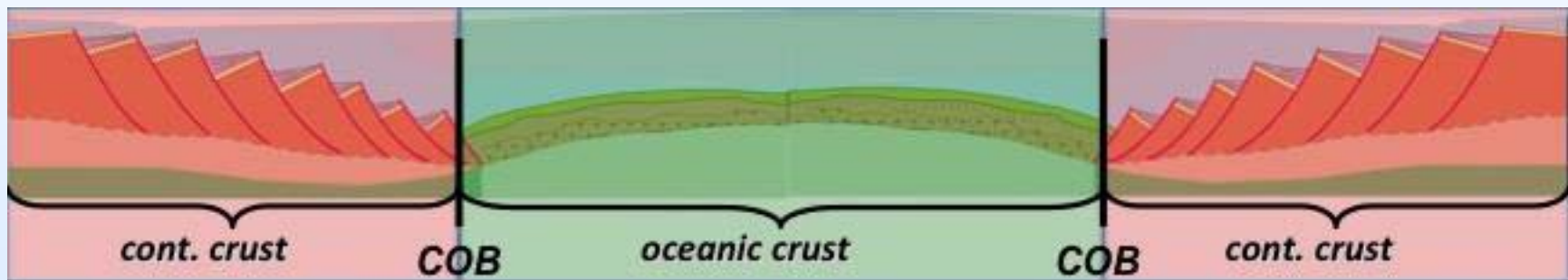
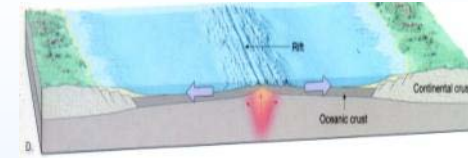
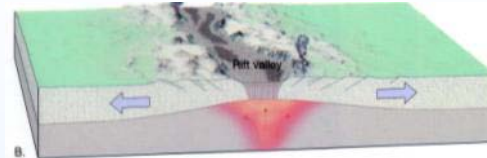
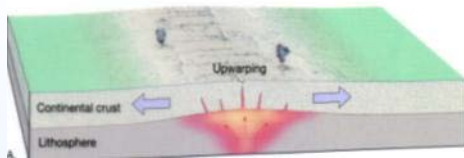
Evolution and architecture of rifted margins "the classical rift model"



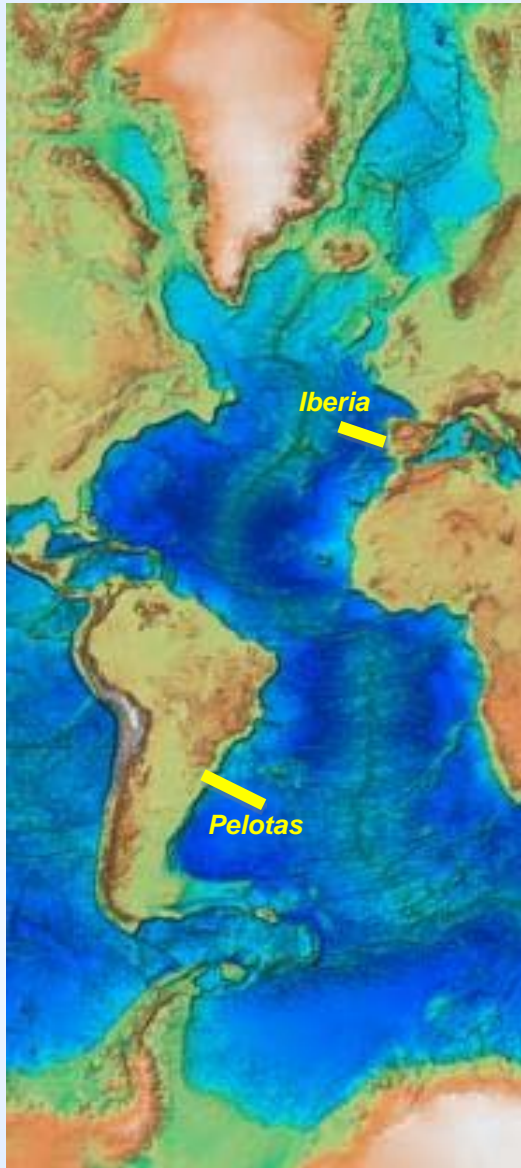
onset of rifting
inherited, equilibrated
continental lithosphere

onset of melting
Decompression melting

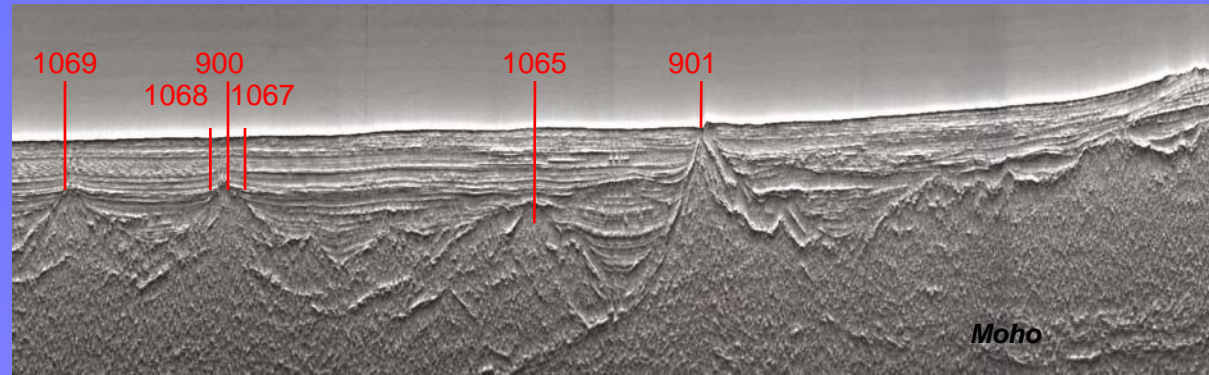
seafloor spreading
controlled by active heat
balance



The “real” view

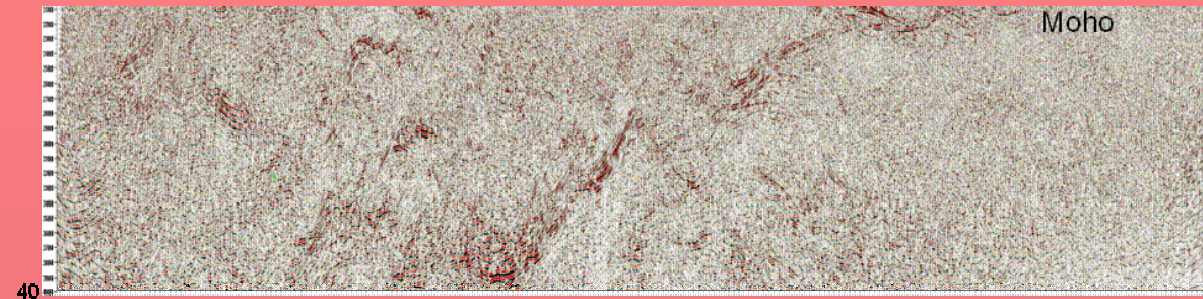


“non”- magmatic



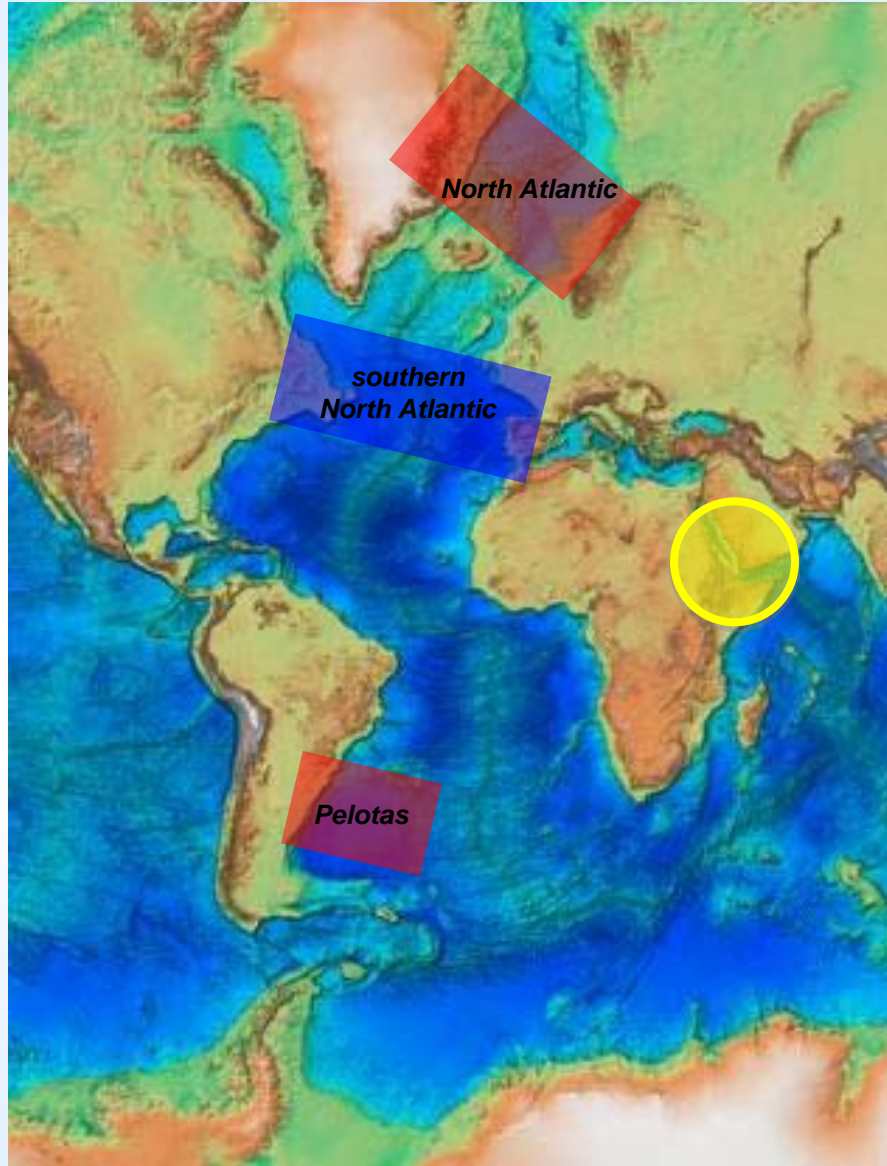
TGS reflection seismic line (migration by Petrobras), with permission of Petrobras

magmatic



PelotasSPAN (ION-GXT) (taken from their homepage)

High-quality reflection seismic data show a divergent style of margin architecture in which quantity and distribution of syn-rift magmatism and fault structures are the most variable features



Major question

do rifted margins operate as either magmatic or non-magmatic or do tectonic and magmatic processes interact before, during and after continental breakup?

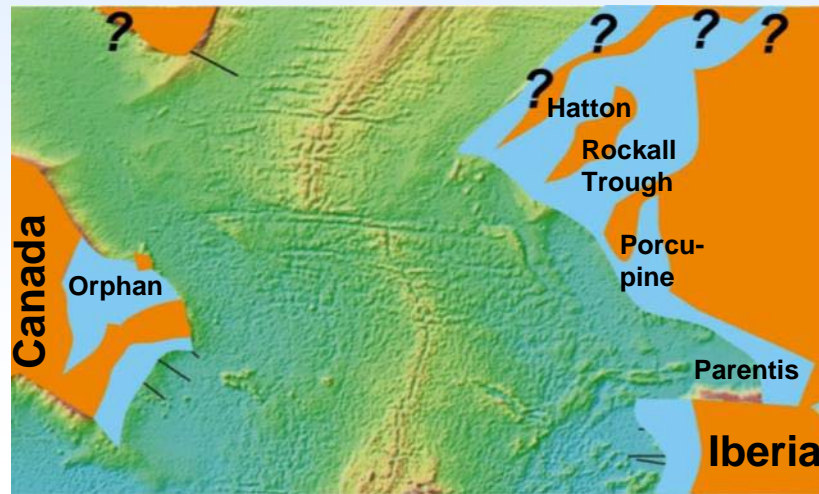
In my presentation

discuss the structural and magmatic evolution of so called “magma-rich” and “magma-poor” rifted margins using examples from the South and North Atlantic

and

I will try to apply the lesson we learned in the Atlantic to the Afar-system

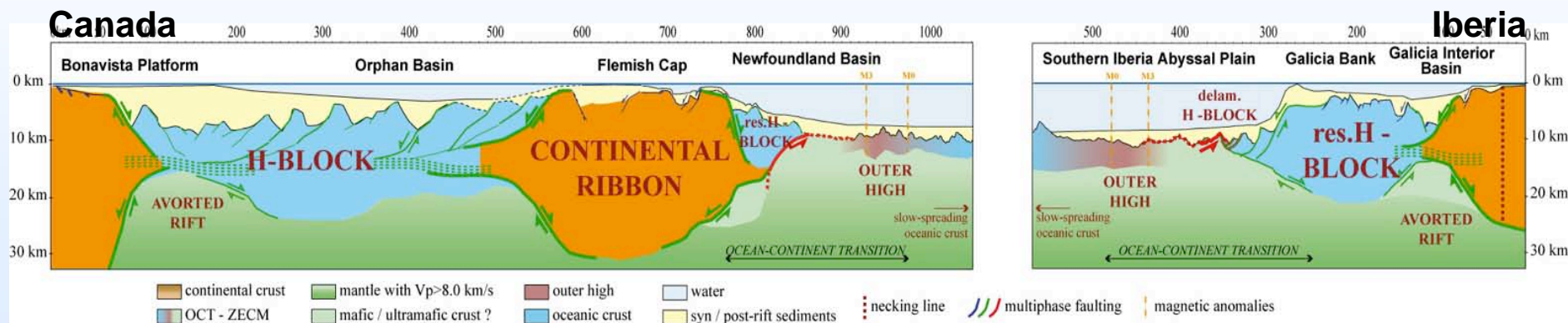
The southern North Atlantic rift-system



Major observations

- Three major crustal domains
- Complex, 3D and polyphase rift structures
- Locale occurrence of aborted rifts floored by lower crust and/or mantle rocks associated with MORB magmatism
- Occurrence of continental ribbons, H-blocks, extensional allochthons and outer highs

Section through the Orphan-Newfoundland-Iberia rift system



modified after Péron-Pinvidic and Manatschal (2010)

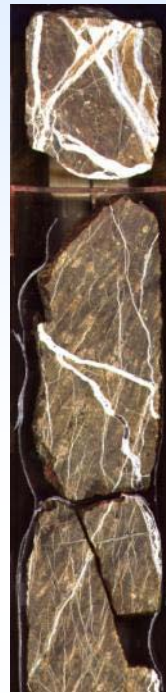
Mantle rocks in the OCT

Drilled mantle lithologies

ODP Leg. 166, 173 and 210

- *Very heterogeneous and strongly serpentinized*
- *Always deformed at the top (cataclasites and gouges)*

**But only highs have been drilled!*



Site 1277



Site 1070

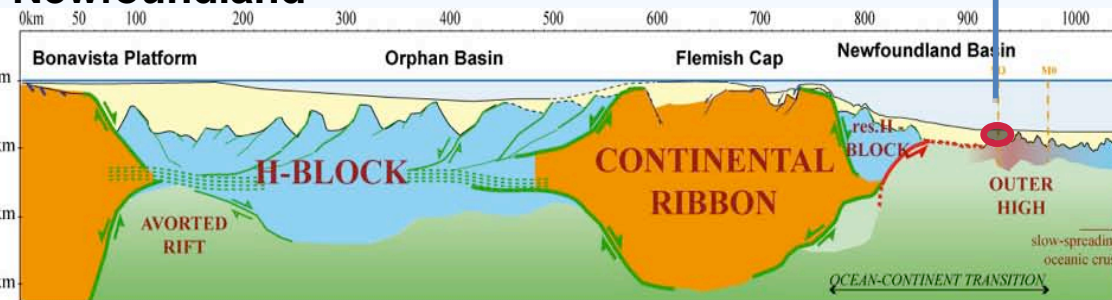


Site 897C

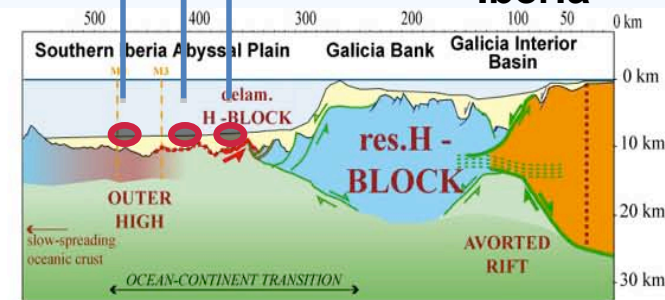


Site 1068

Newfoundland



Iberia

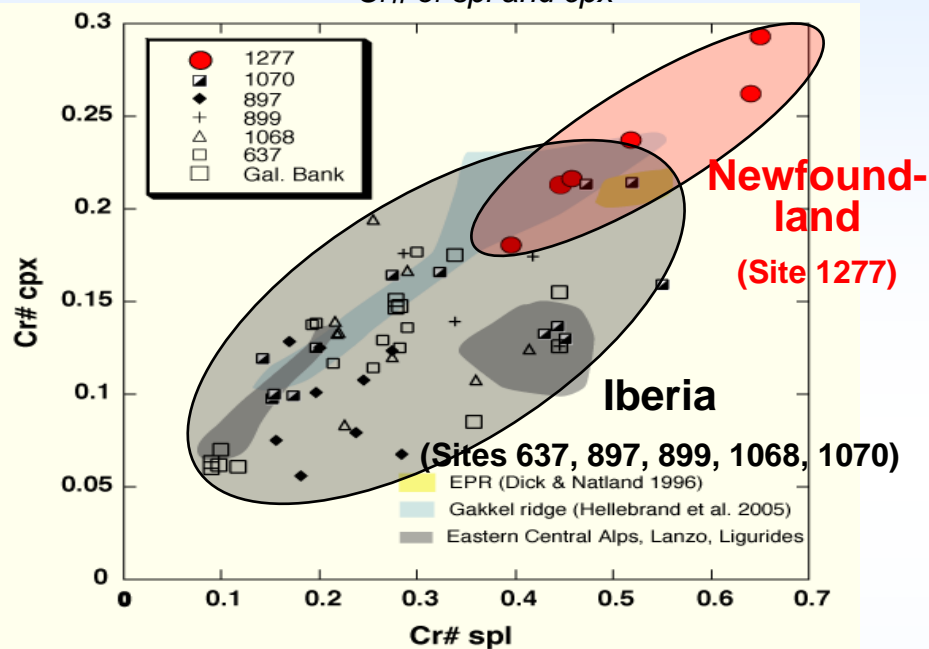


continental crust
 mantle with $V_p > 8.0$ km/s
 outer high
 water
 OCT - ZECM
 mafic / ultramafic crust ?
 oceanic crust
 syn / post-rift sediments
 necking line
 multiphase faulting
 magnetic anomalies

Composition of the mantle rocks

Chemistry of mantle rocks

Cr# of spl and cpx

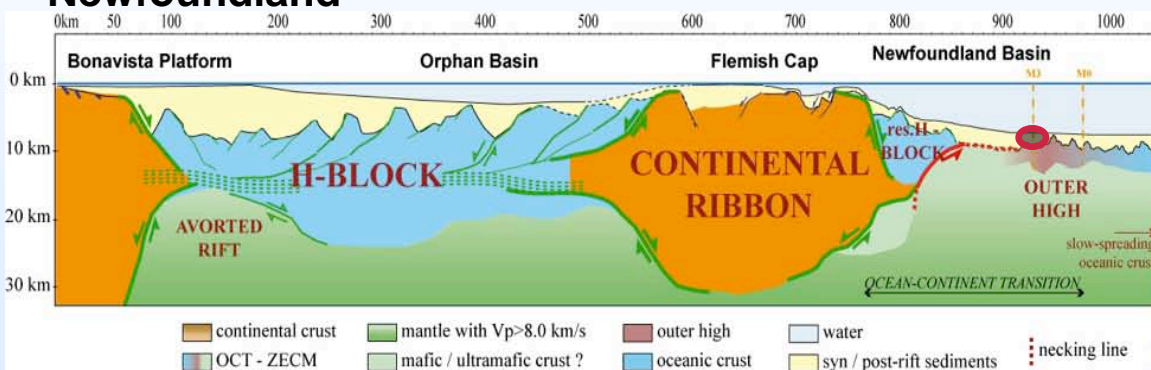


Müntener and Manatschal, 2006

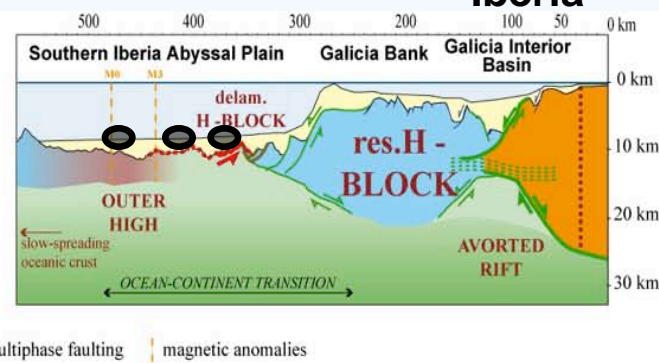
Major conclusions

- Newfoundland peridotite more depleted than Iberia (*higher degrees of melting*)
- All holes from Iberia show some deviations from Cr# correlation (*equilibration with plagioclase*)
- Newfoundland peridotite among most depleted peridotites drilled (*inherited mantle domain*)
- Some similarities with sub-arc mantle (*imprint of old subduction?*)

Newfoundland



Iberia

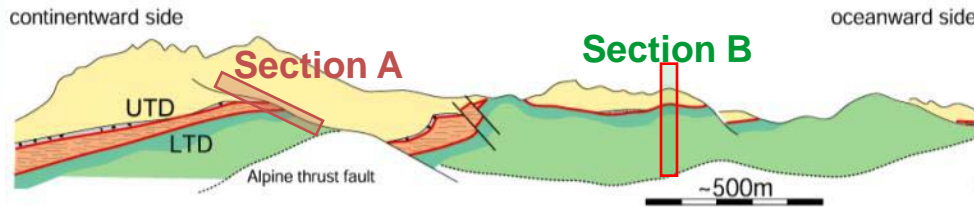


modified after Péron-Pinvidic and Manatschal (2010)

Example of exhumed mantle in the Alps

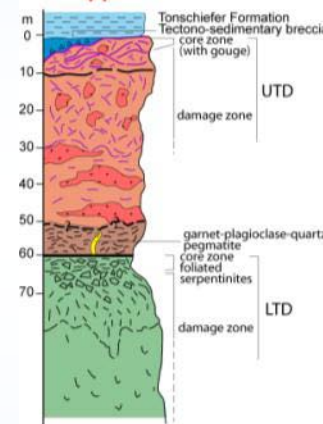


Tasna Ocean Continent Transition (OCT)

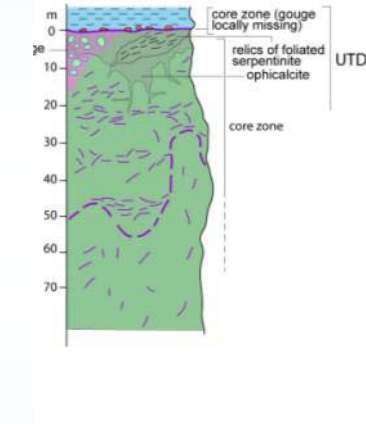


Florineth and Frotzheim 1994, Manatschal et al. 2006

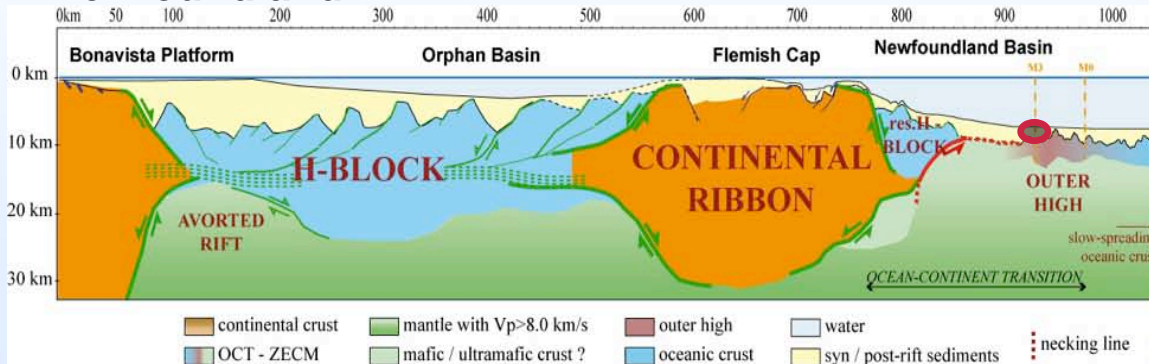
Section A



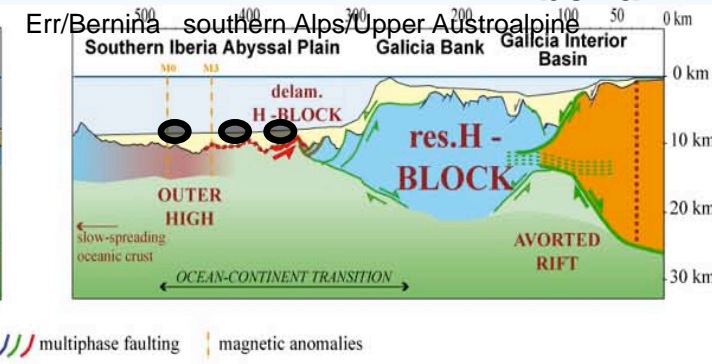
Section B



Newfoundland



Iberia

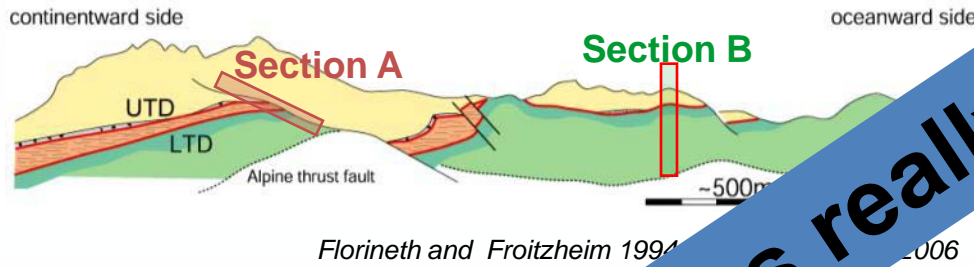


modified after Péron-Pinvidic and Manatschal (2010)

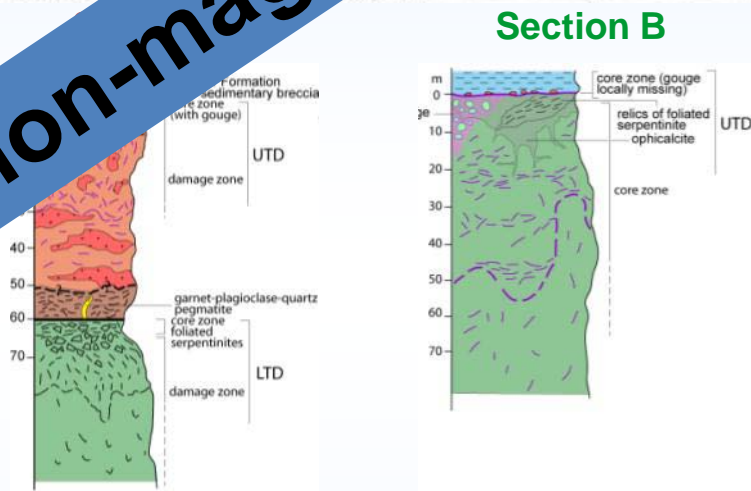
Example of exhumed mantle in the Alps



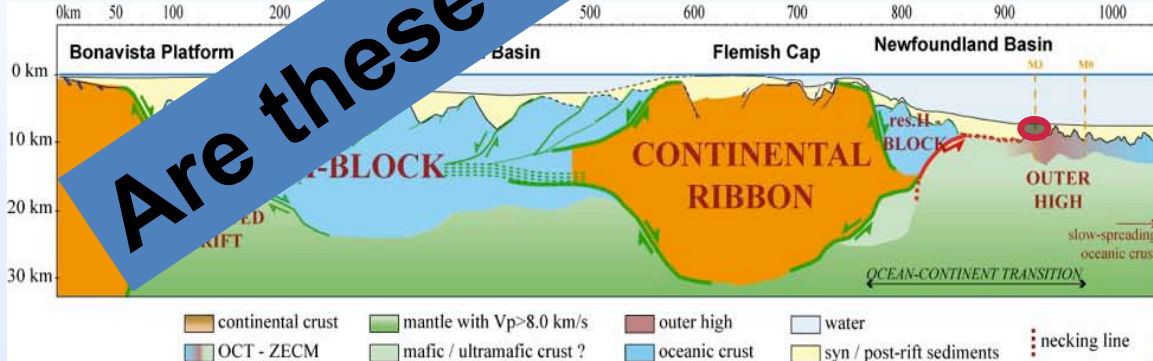
Tasna Ocean Continent Transition (OCT)



Florineth and Frotzheim 1994, 2006

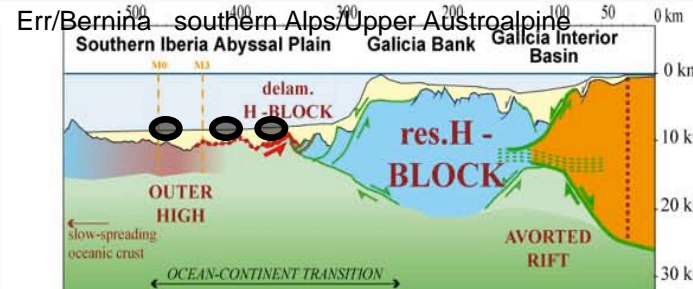


Newfoundland



continental crust
 mantle with $V_p > 8.0$ km/s
 outer high
 water
 OCT - ZECM
 mafic / ultramafic crust ?
 oceanic crust
 syn / post-rift sediments
 necking line
 multiphase faulting
 magnetic anomalies

Iberia



modified after Péron-Pinvidic and Manatschal (2010)

Are these systems really non-magmatic ?

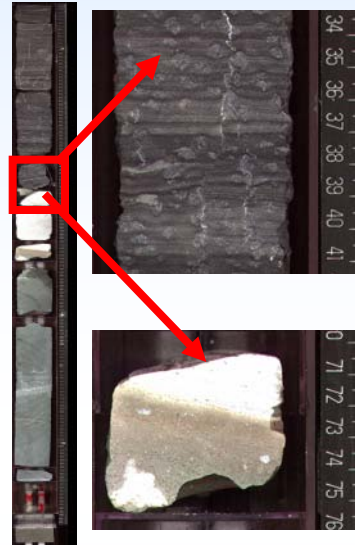
Magmatic rocks in the OCT (*Iberia-Newfoundland and Alpine Tethys*)

Drilled magmatic rocks

(ODP Sites 897, 1070, 1276, 1277)

- Magmatic rocks are present but rare
- Syn-rift as well as post-rift magmatic activity
- Apparent increase in magma oceanwards
- Alkaline as well as MOR signatures

Sill in sediments



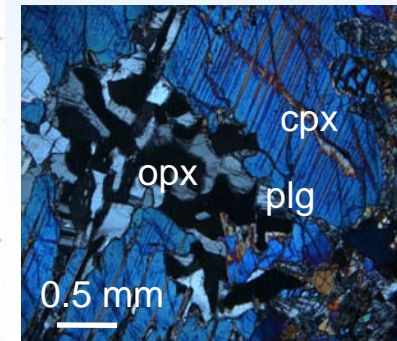
Core 1276-87-6

Gabbros in exhumed mantle



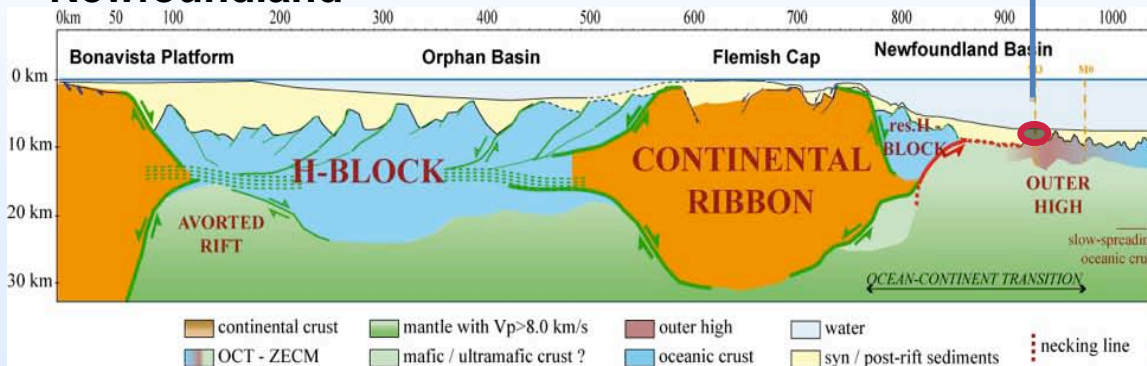
Site 1070

Infiltrated mantle

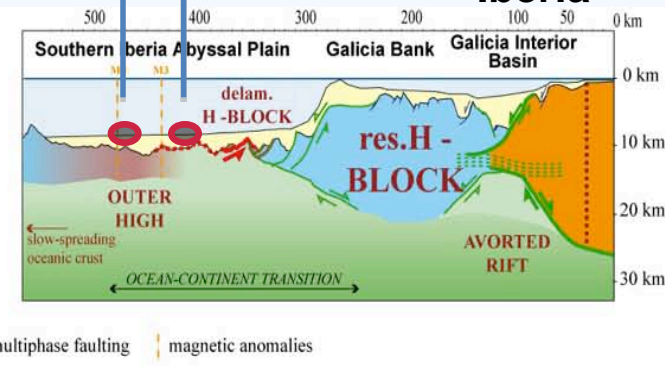


Site 897 4R-1

Newfoundland



Iberia

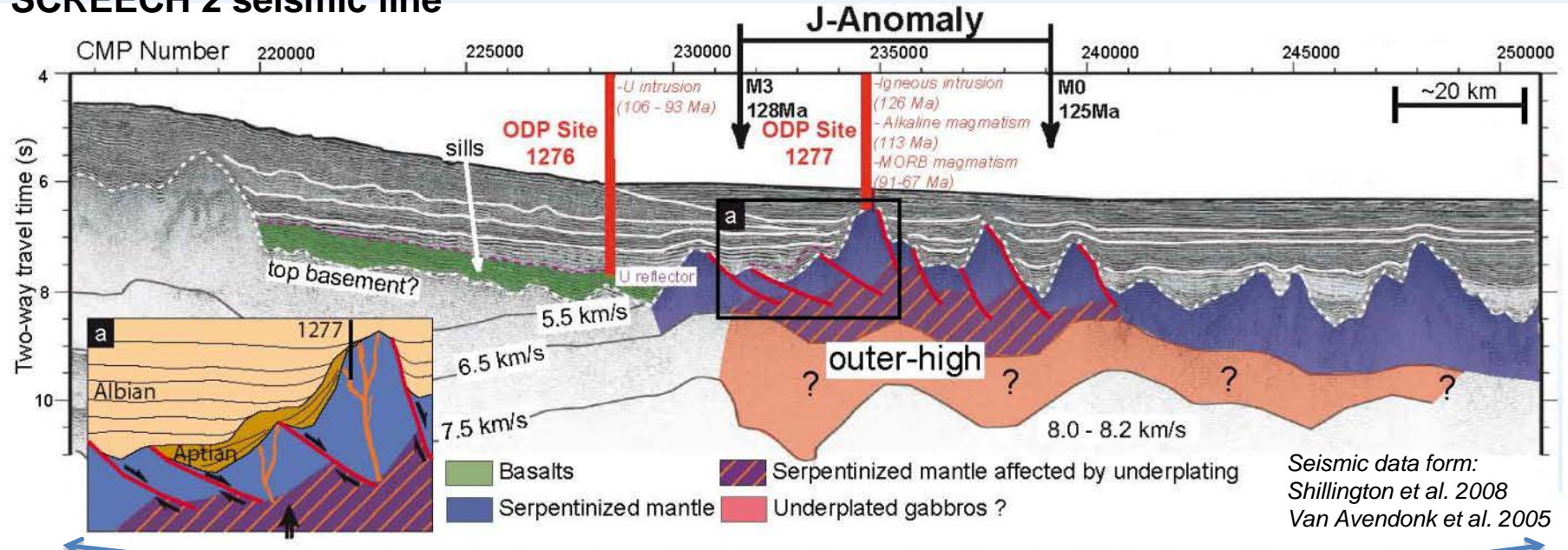


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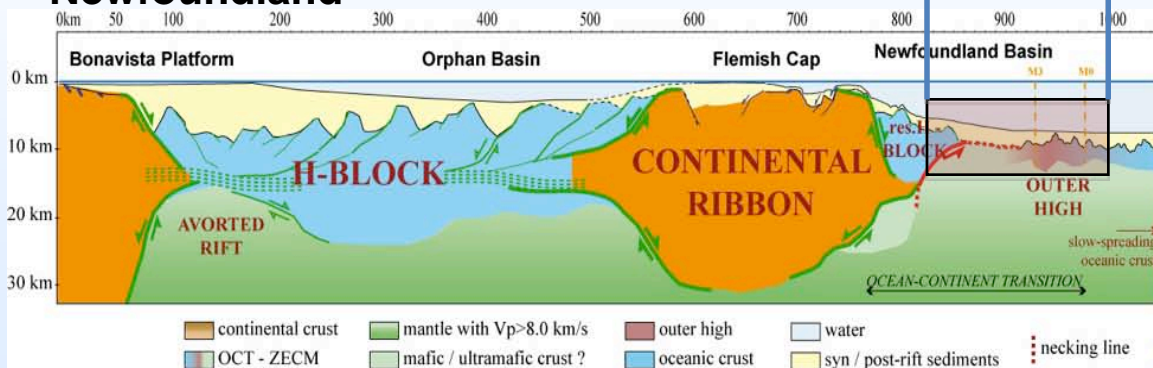
Where and when did breakup occur?

Bronner et al. 2011 (Nature Geoscience)

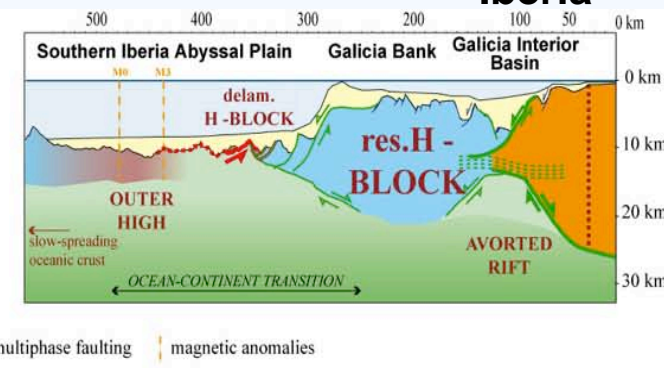
SCREECH 2 seismic line



Newfoundland



Iberia

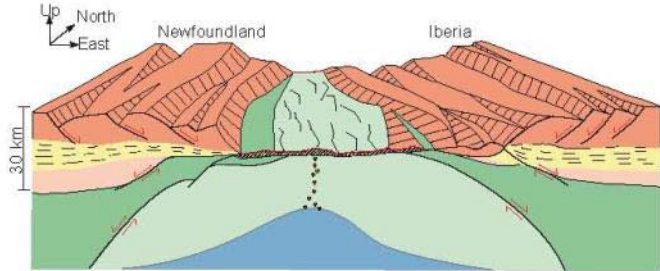


modified after Péron-Pinvidic and Manatschal (2010)

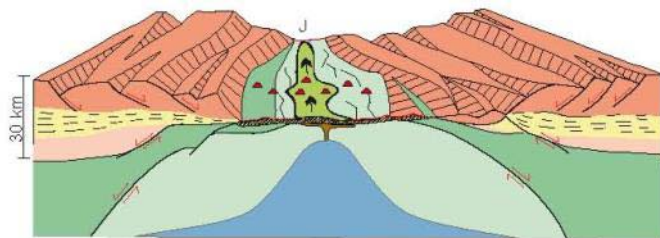
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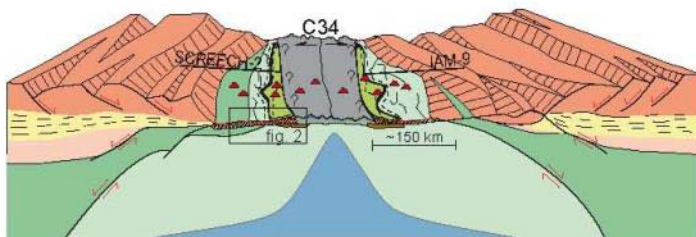
a Mantle exhumation Valanginian to late Aptian (140 to 112 Ma)



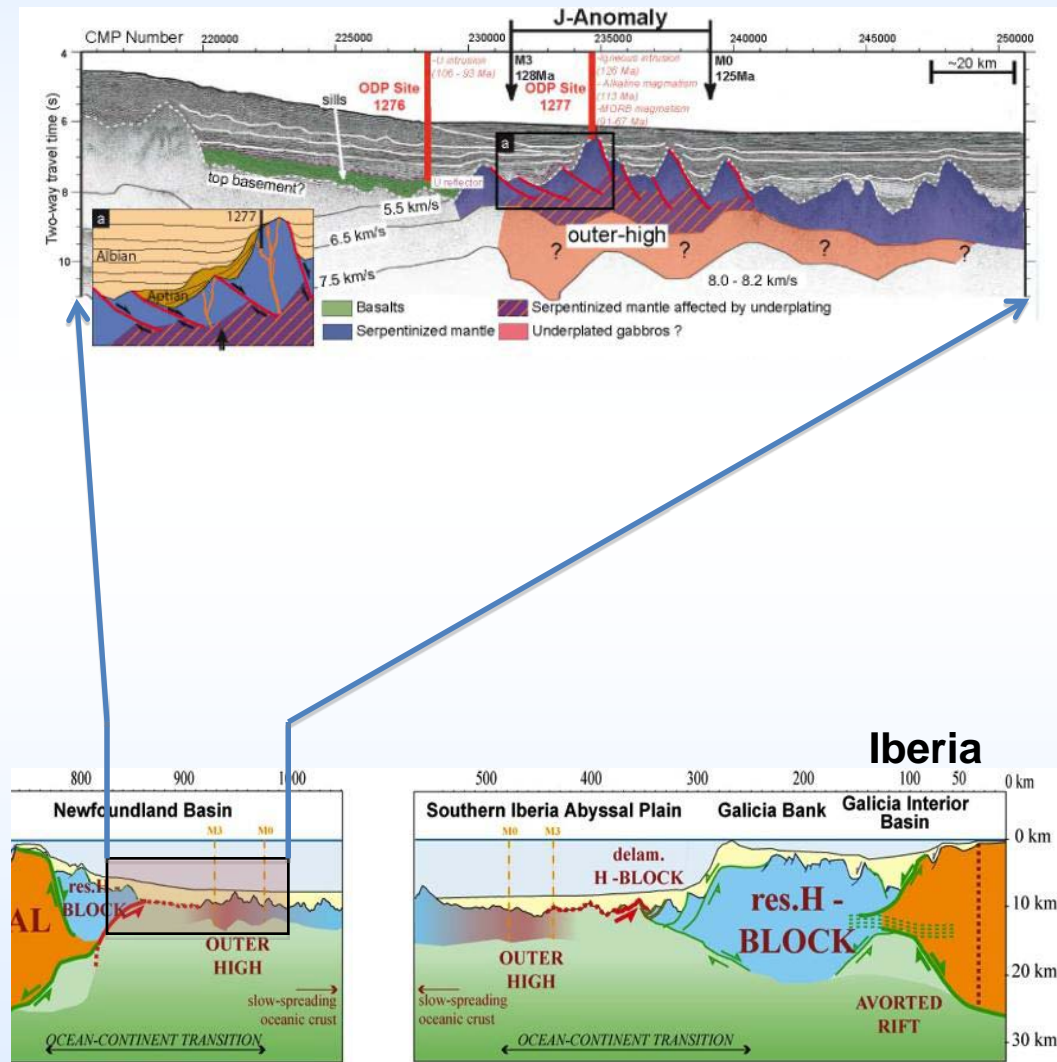
b Breakup Aptian/Albian (~112 Ma)



c Early seafloor spreading Albian to Campanian (112 - 84 Ma)



- Upper continental crust
- Middle - lower continental crust
- Inherited subcontinental mantle
- Infiltrated mantle
- Serpentinized mantle
- Asthenospheric mantle
- Basalt
- Magmatic intrusion
- Gabbro
- Slow spreading oceanic crust
- Serpentinized mantle affected by underplating
- Underplated gabbros ?
- J anomaly
- Seamount



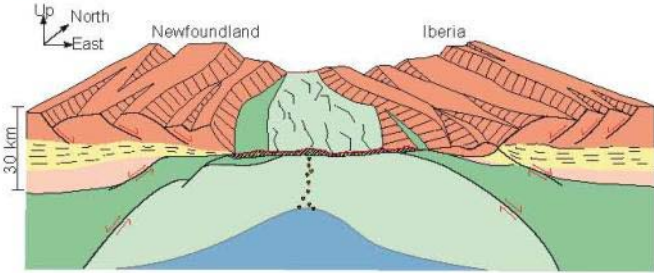
water
 yn / post-rift sediments necking line multiphase faulting magnetic anomalies

modified after Péron-Pinvidic and Manatschal (2010)

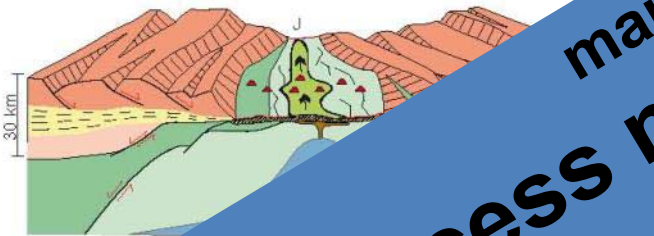
Where and when did breakup occur?

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a Mantle exhumation Valanginian to late Aptian (140 to 112 Ma)



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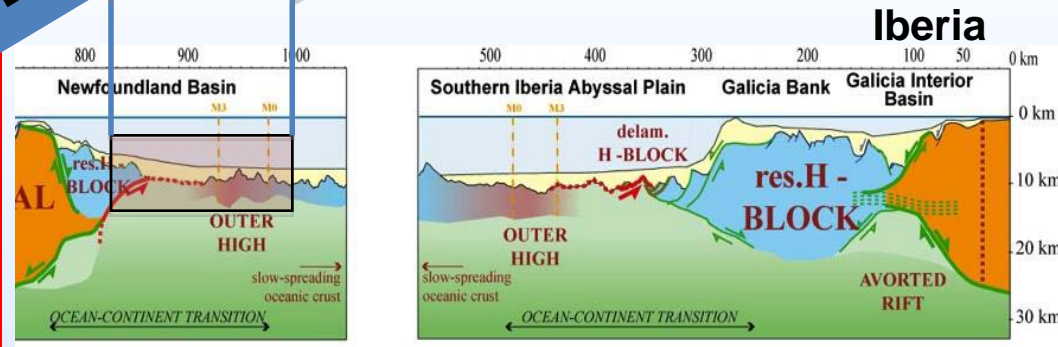
c Early



- Upper continental crust
- Middle continental crust
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- Infiltrated mantle
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- Seamount
- J anomaly
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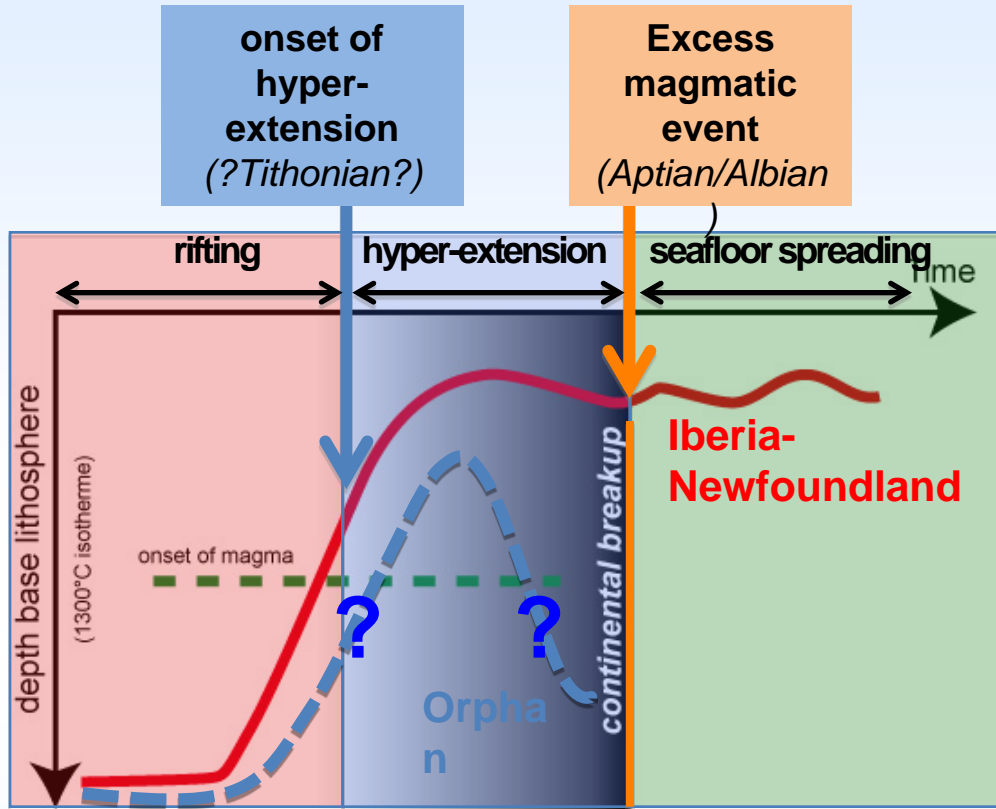


major conclusion:
 an excess magmatic event at the
 Aptian-Albian boundary is at the origin
 of continental breakup



- water
- syn / post-rift sediments
- necking line
- multiphase faulting
- magnetic anomalies

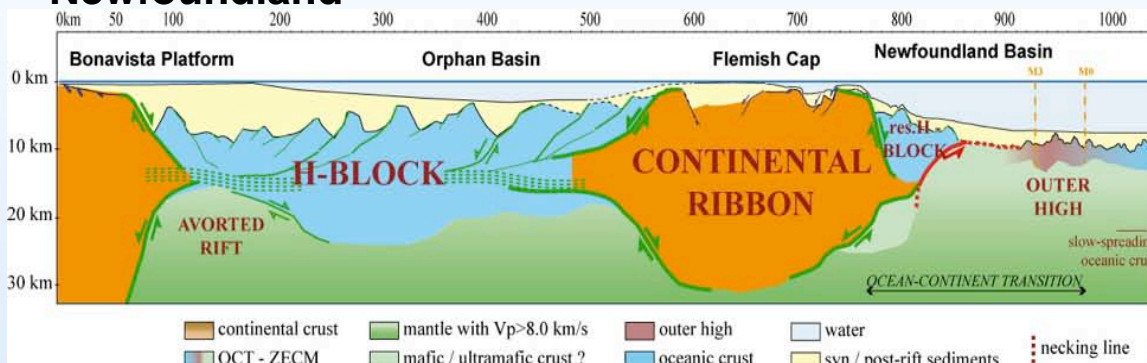
The southern North Atlantic: conclusions



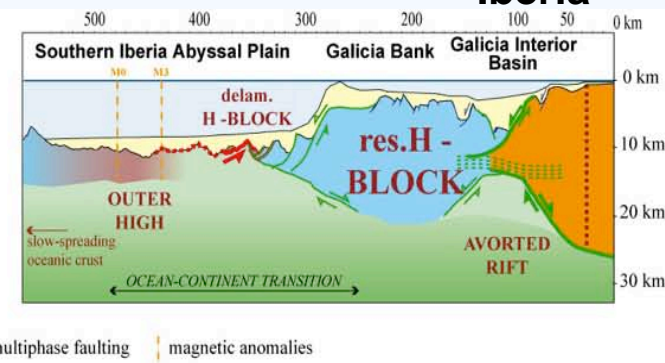
Major conclusions:

- Magma is of subordinate importance during hyper-extension
- Breakup is linked with excess magmatic event
- More than 150 km of exhumed mantle

Newfoundland



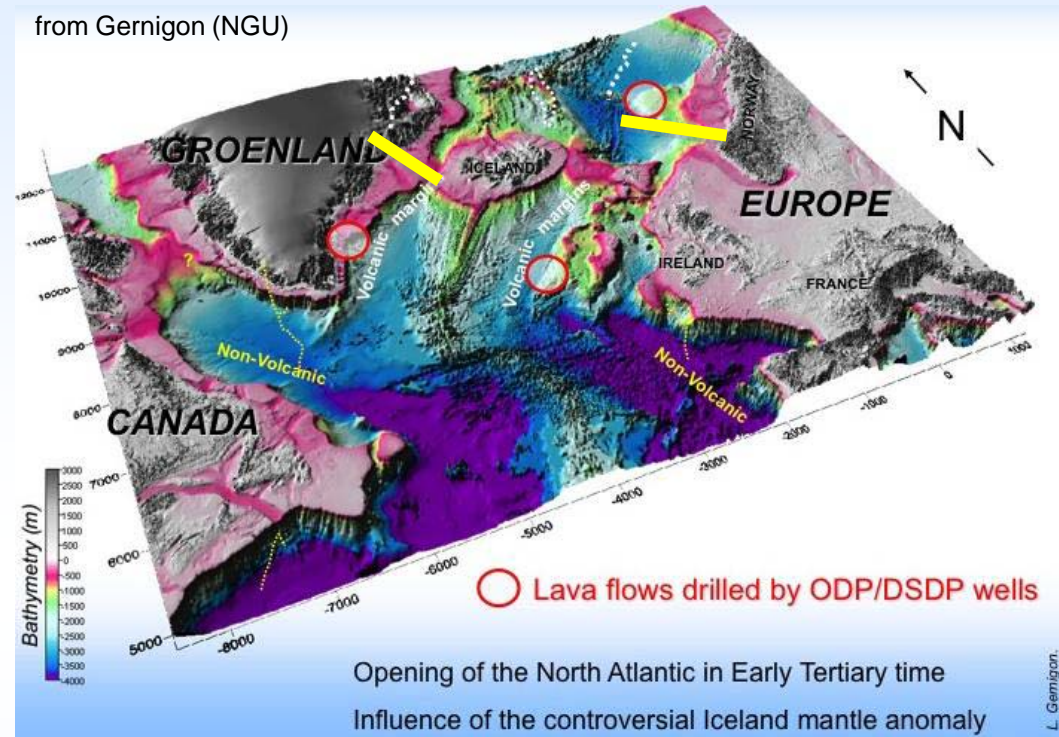
Iberia



The North Atlantic rift-system

Major observations

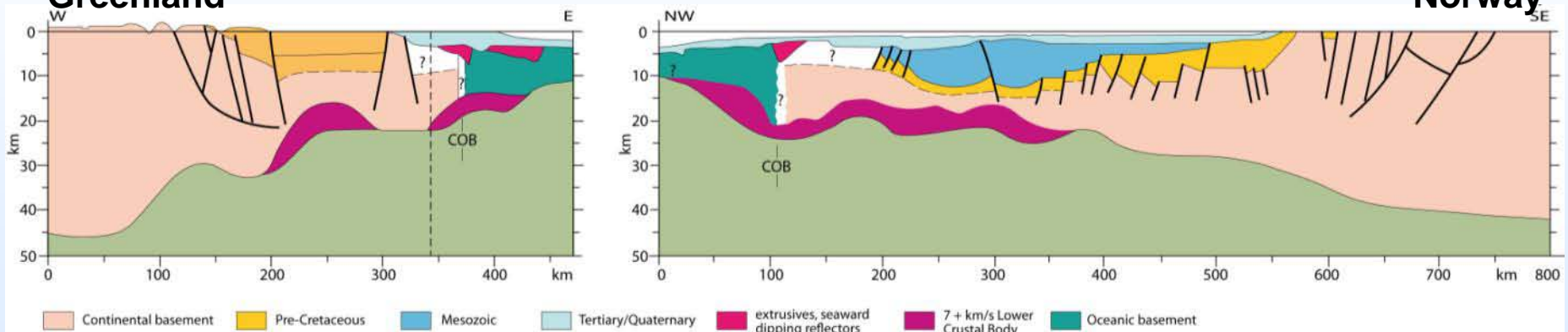
- Prominent magmatic event during continental breakup related to the arrival of the Iceland plume
- Present-day rift structure is dominated by magmatic breakup



Section through the East Greenland – mid-Norway conjugate margins

Greenland

Norway

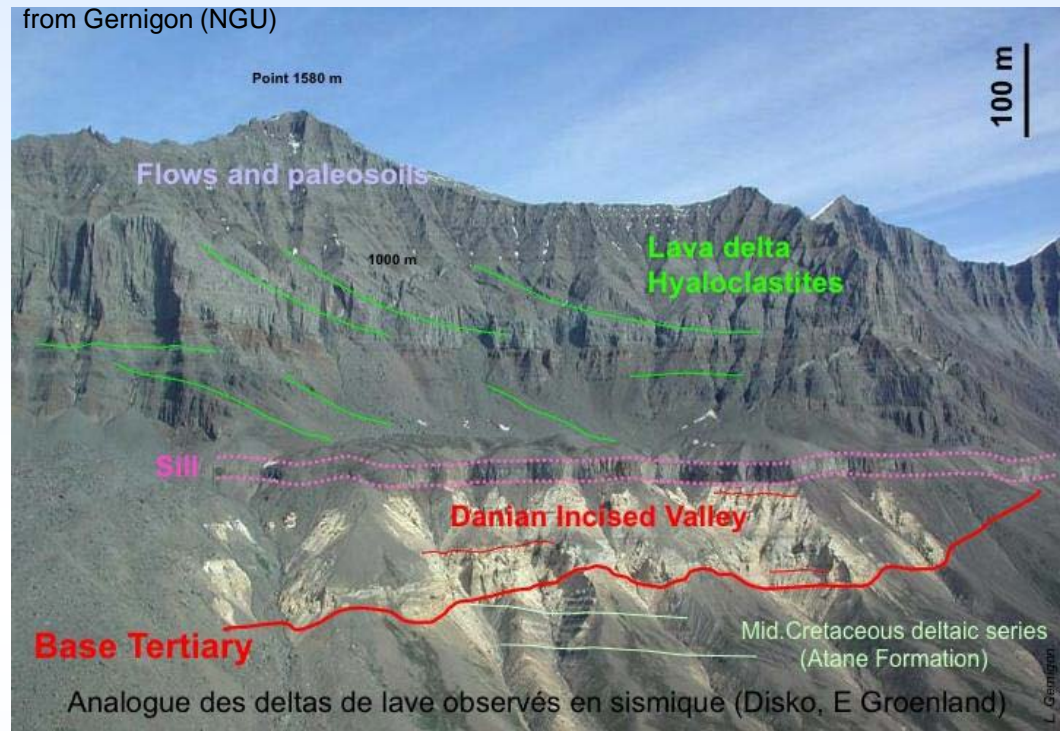


(modified from Tsikalas et al. 2007)

The North Atlantic rift-system

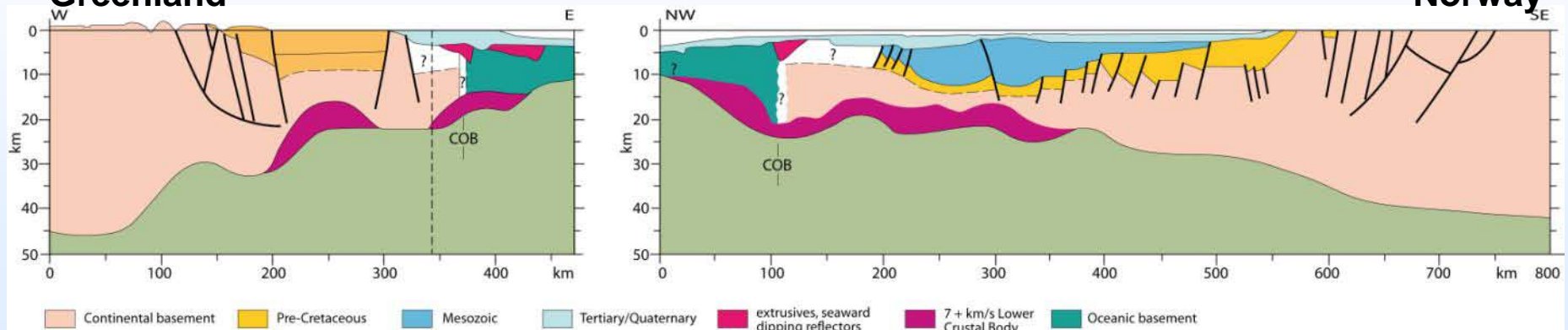
Major observations

- Prominent magmatic event during continental breakup related to the arrival of the Island plume
- Present-day rift structure is dominated by magmatic breakup



Greenland

Norway



(modified from Tsikalas et al. 2007)

The North Atlantic rift-system

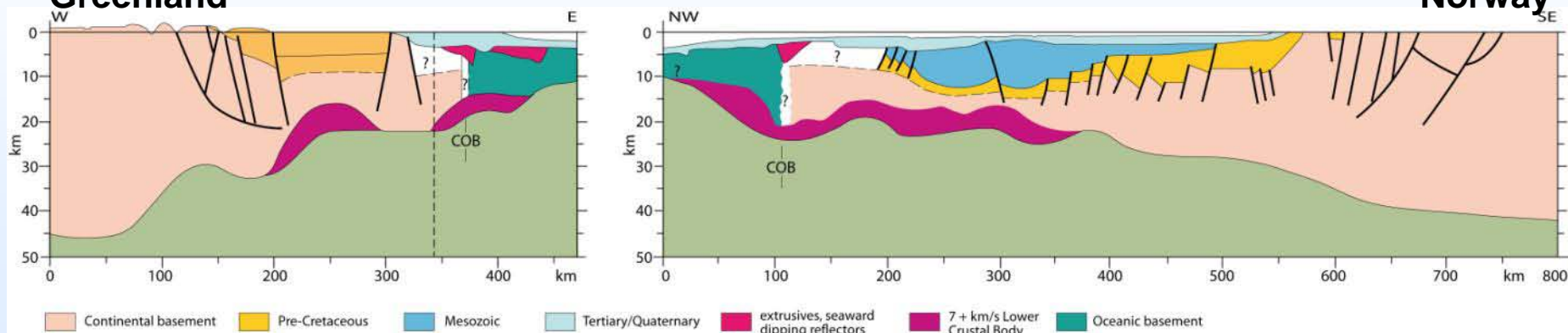
Major observations

- Prominent magmatic event during continental breakup related to the arrival of the Island plume
- Present-day rift structure is dominated by magmatic breakup



Greenland

Norway

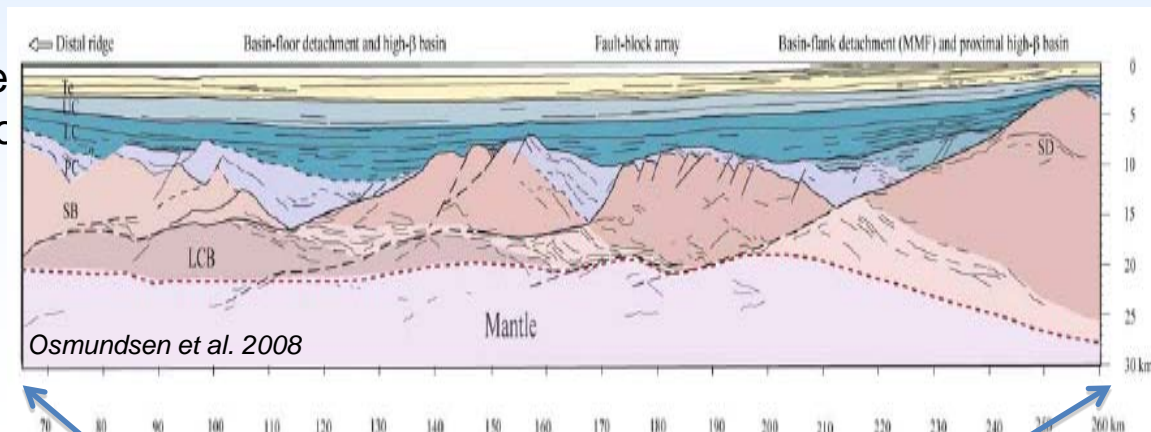


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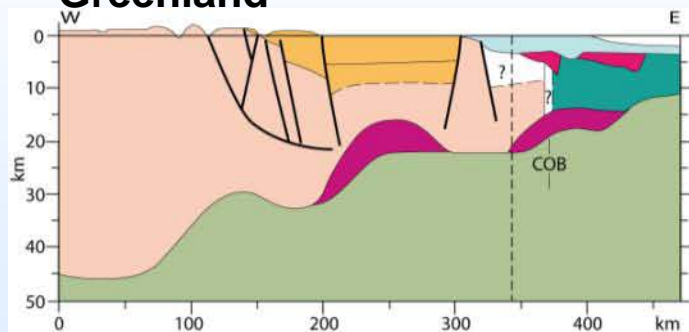
The North Atlantic rift-system

Major observations

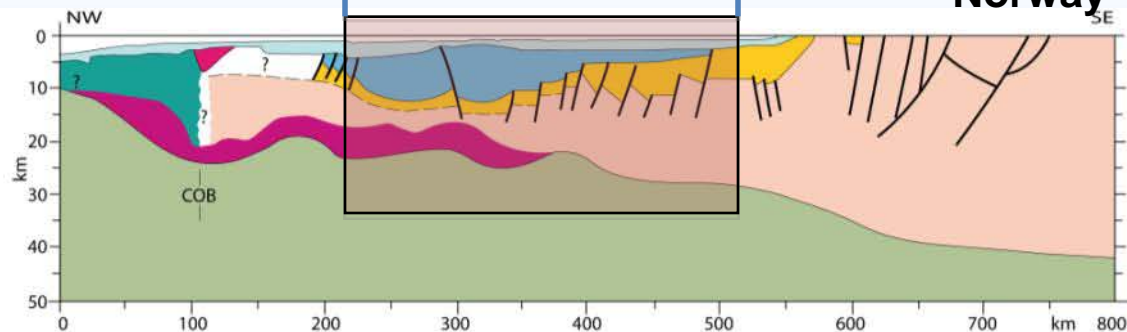
- Late Jurassic – Early Cretaceous hyper-extension resulted in extreme thinning of the continental crust in the present-day Møre and Vøring basins
- Excess magmatic event overprints hyper-extended continental crust



Greenland



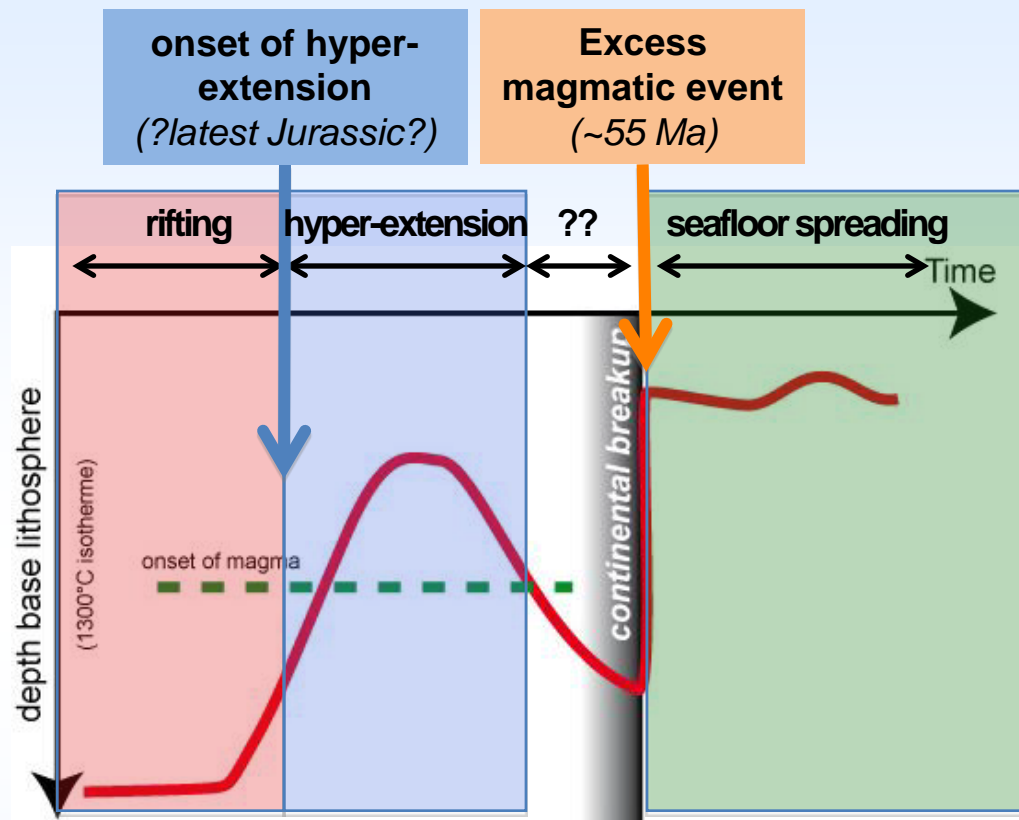
Norway



Continental basement
 Pre-Cretaceous
 Mesozoic
 Tertiary/Quaternary
 extrusives, seaward dipping reflectors
 7 + km/s Lower Crustal Body
 Oceanic basement

(modified from Tsikalas et al. 2007)

North Atlantic rift-system: conclusions



Major observations

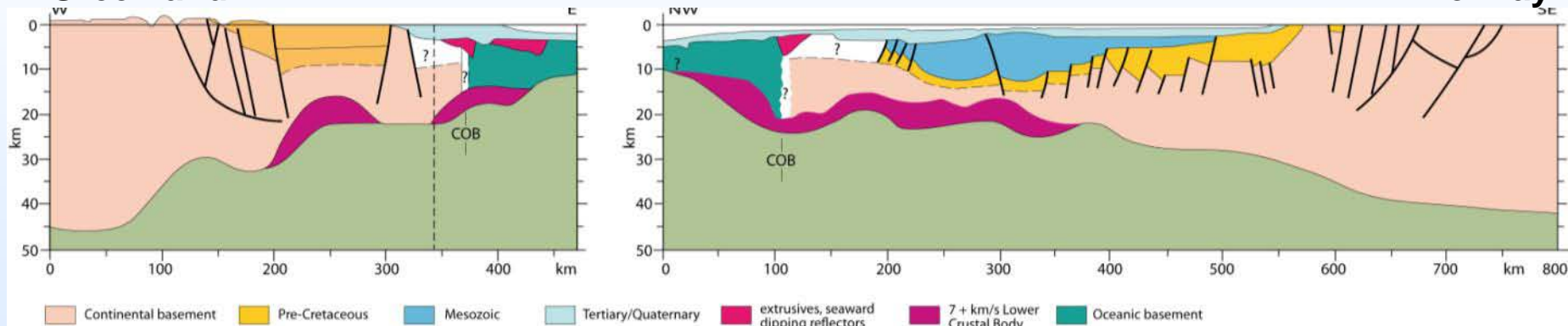
- Late Jurassic – Early Cretaceous hyper-extension
- Breakup (~55 Ma) related to excess magmatic event (SDRs and Low Crustal Magmatic Bodies)

Major question

- Why did the Late Jurassic – Early Cretaceous event not form more magma, and why did the North Atlantic not open during this stage?

Greenland

Norway



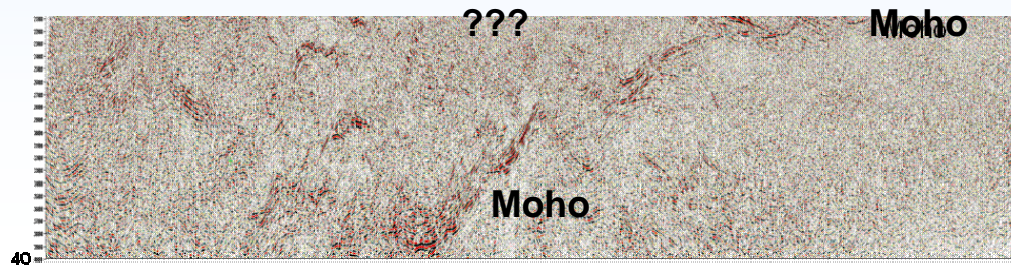
(modified from Tsikalas et al. 2007)

The Pelotas rift system (S-Brazil)

Major observations

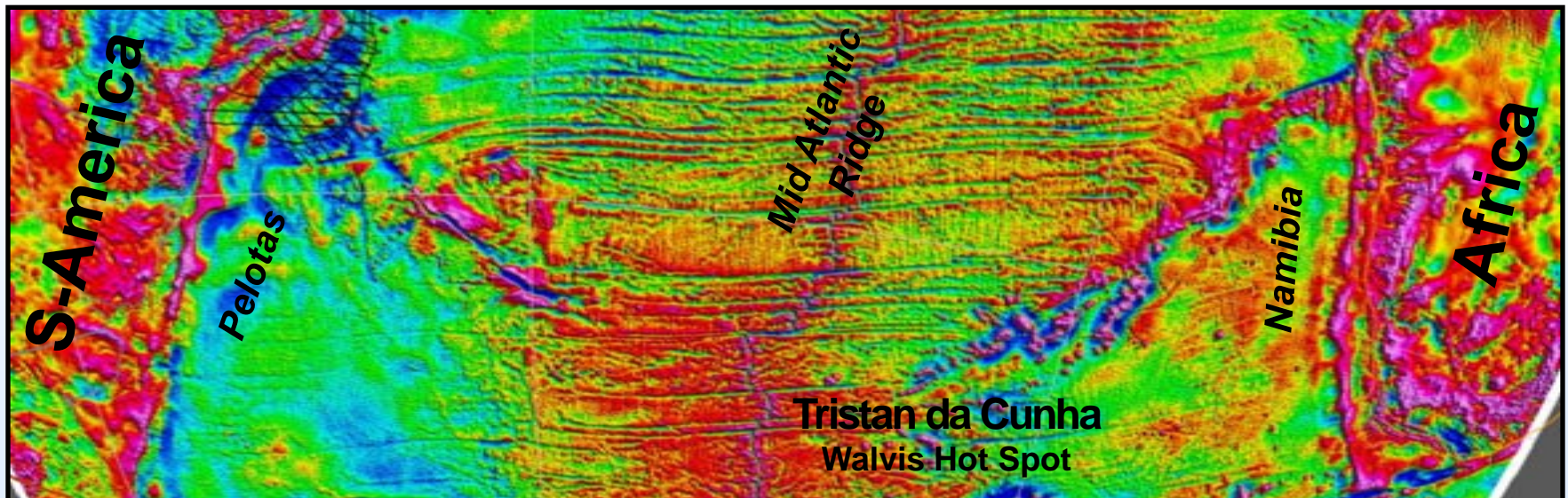
- Well developed SDR sequences
- Strong Moho-reflections
- What is underneath the SDR?

SDR

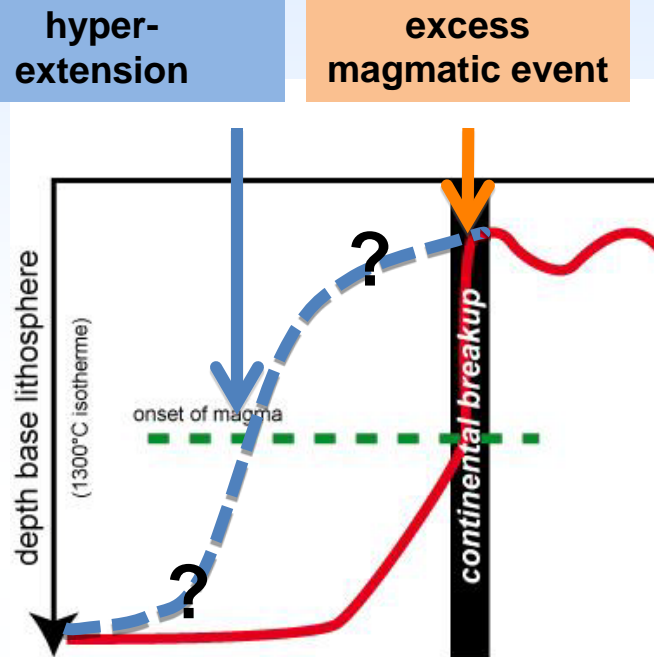


PelotasSPAN (ION-GXT) (taken from their homepage)

Gravimetric Free-air Map *(Sandwell and Smith (2009))*



South Atlantic (Pelotas – S-Brazil)



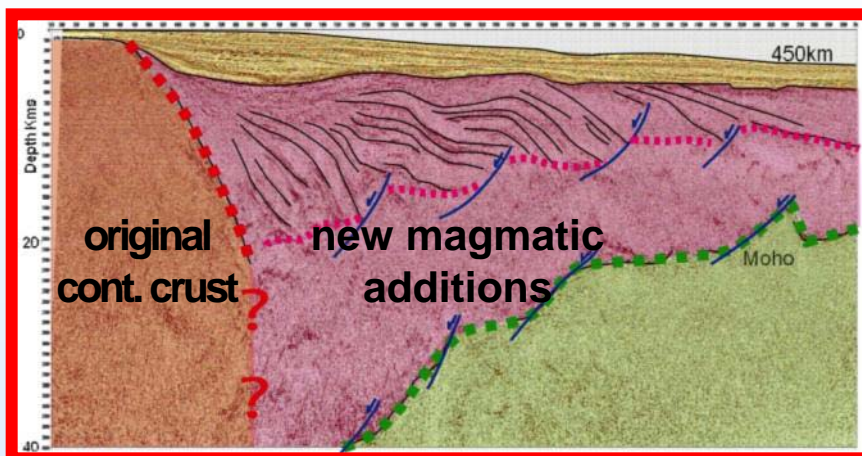
Interpretation I

- *Whole crust is formed by newly formed magmatic crust*

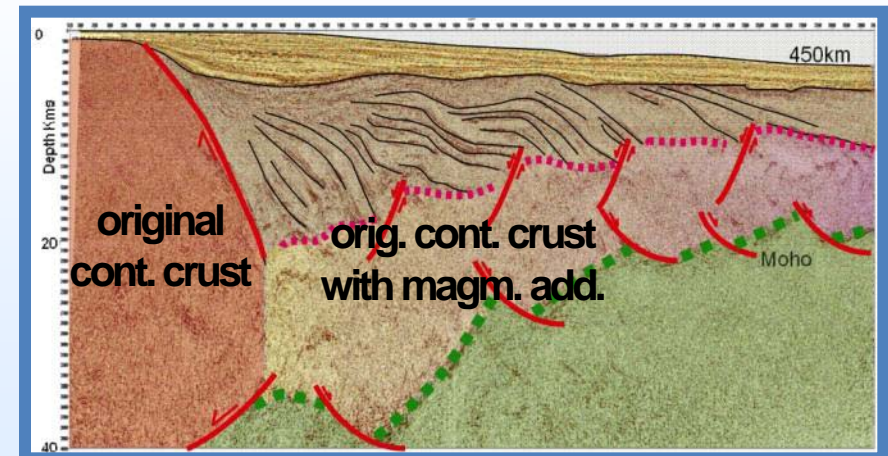
Interpretation II

- *Rifting initiates by hyper-extension*
- *Breakup is linked with excess magmatic event overprinting hyper-extended crust*

Purely magmatic

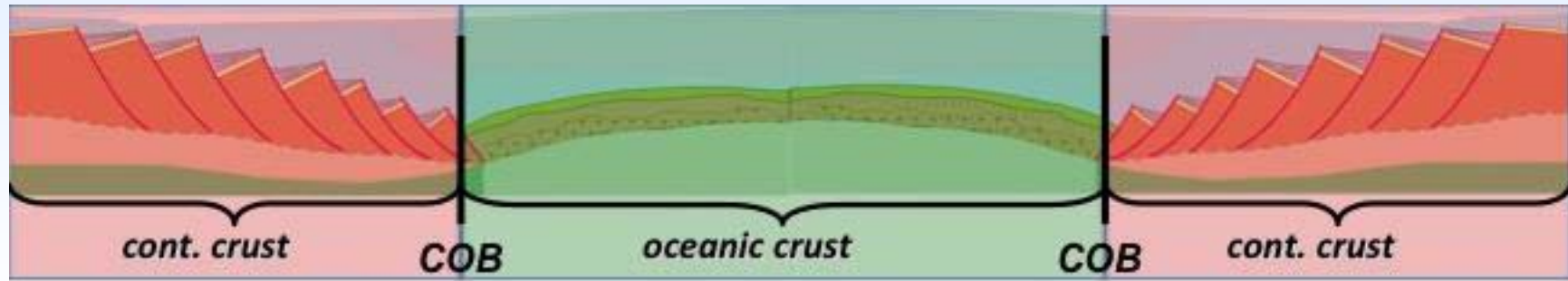


Hyper-extension with magmatic additions



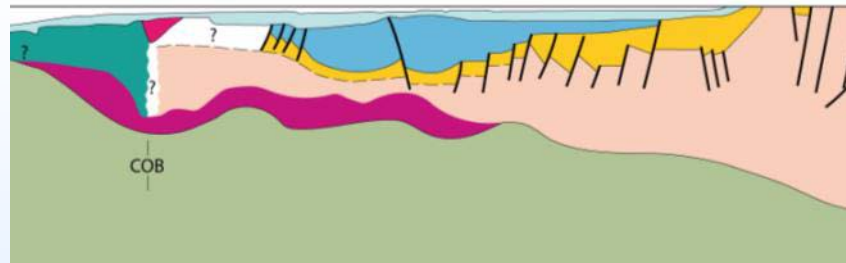
How do continents breakup?

The classical rift model



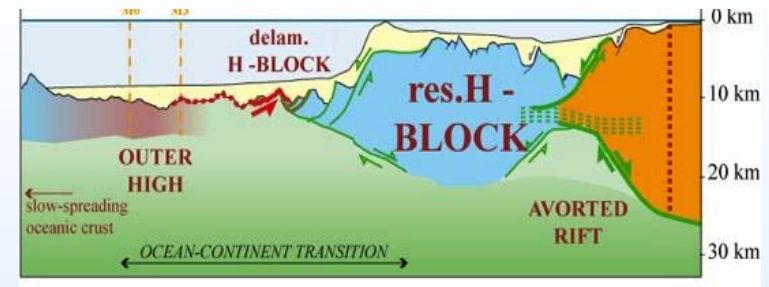
Results of the Ocean Drilling Program and seismic imaging

magmatic margins



- Seaward Dipping Reflection Sequences (SDR)
- Thickened crust (magmatic additions)

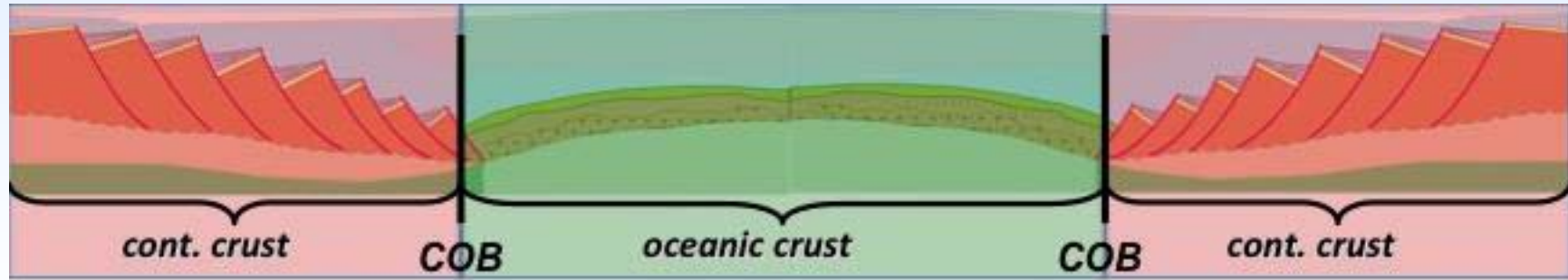
“non”-magmatic margins



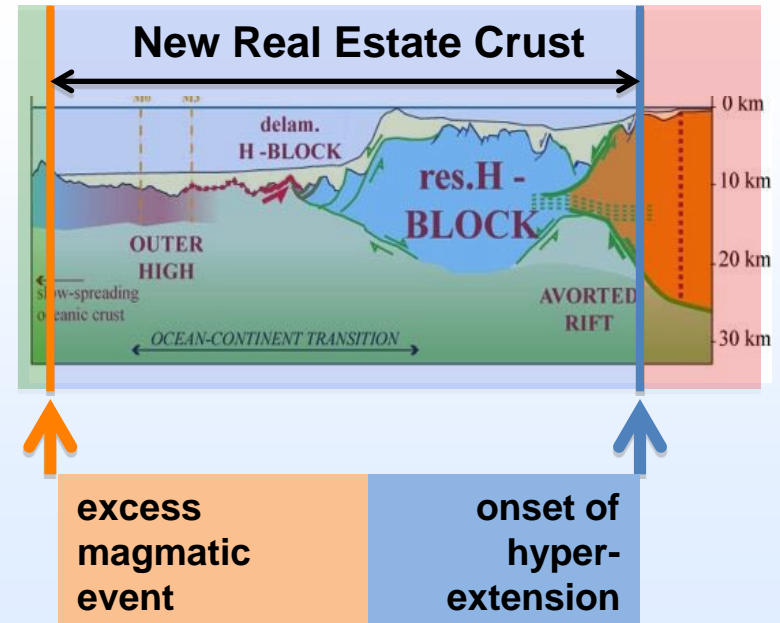
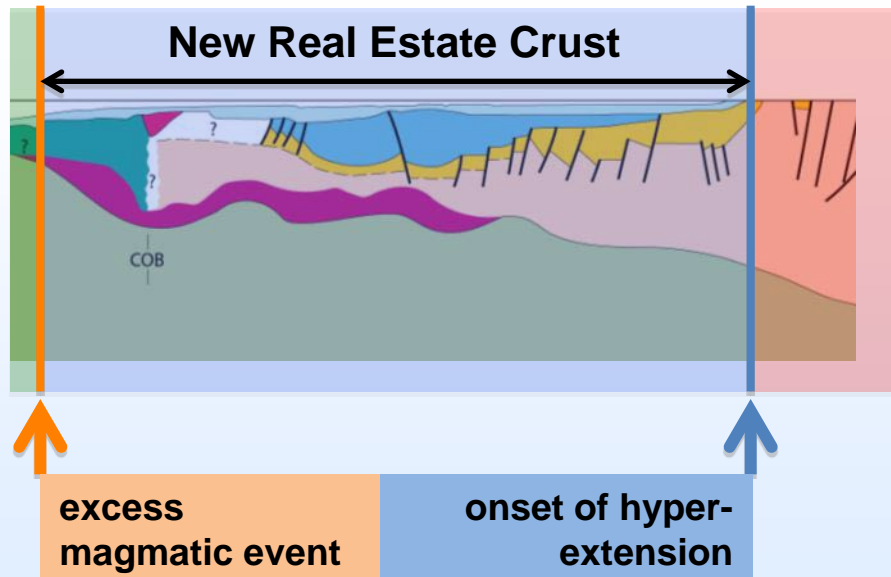
- Exhumed mantle
- Necking zones
- Detachment faulting and extensional allochthons

How do ocean breakup?

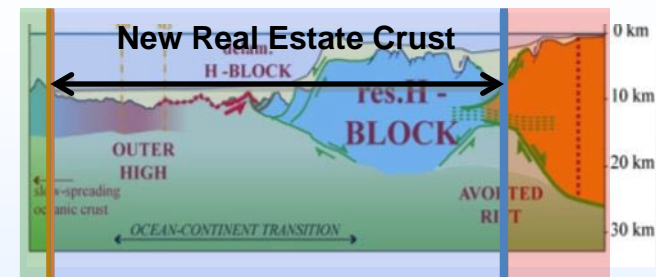
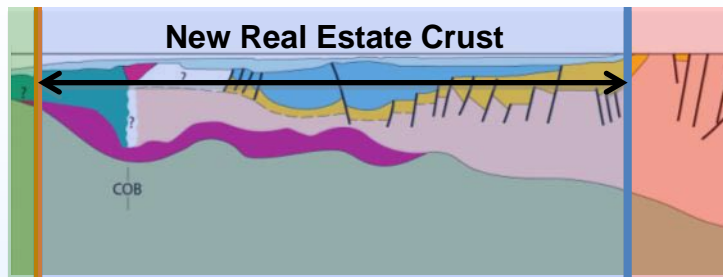
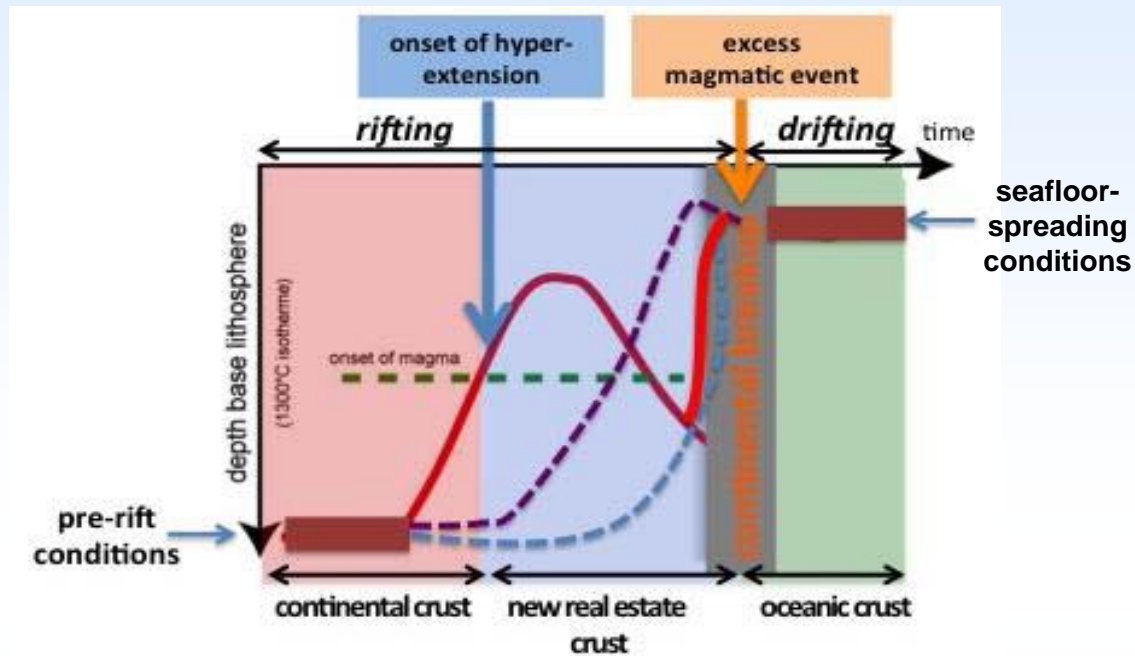
The classical rift model



another approach to describe rifted margins



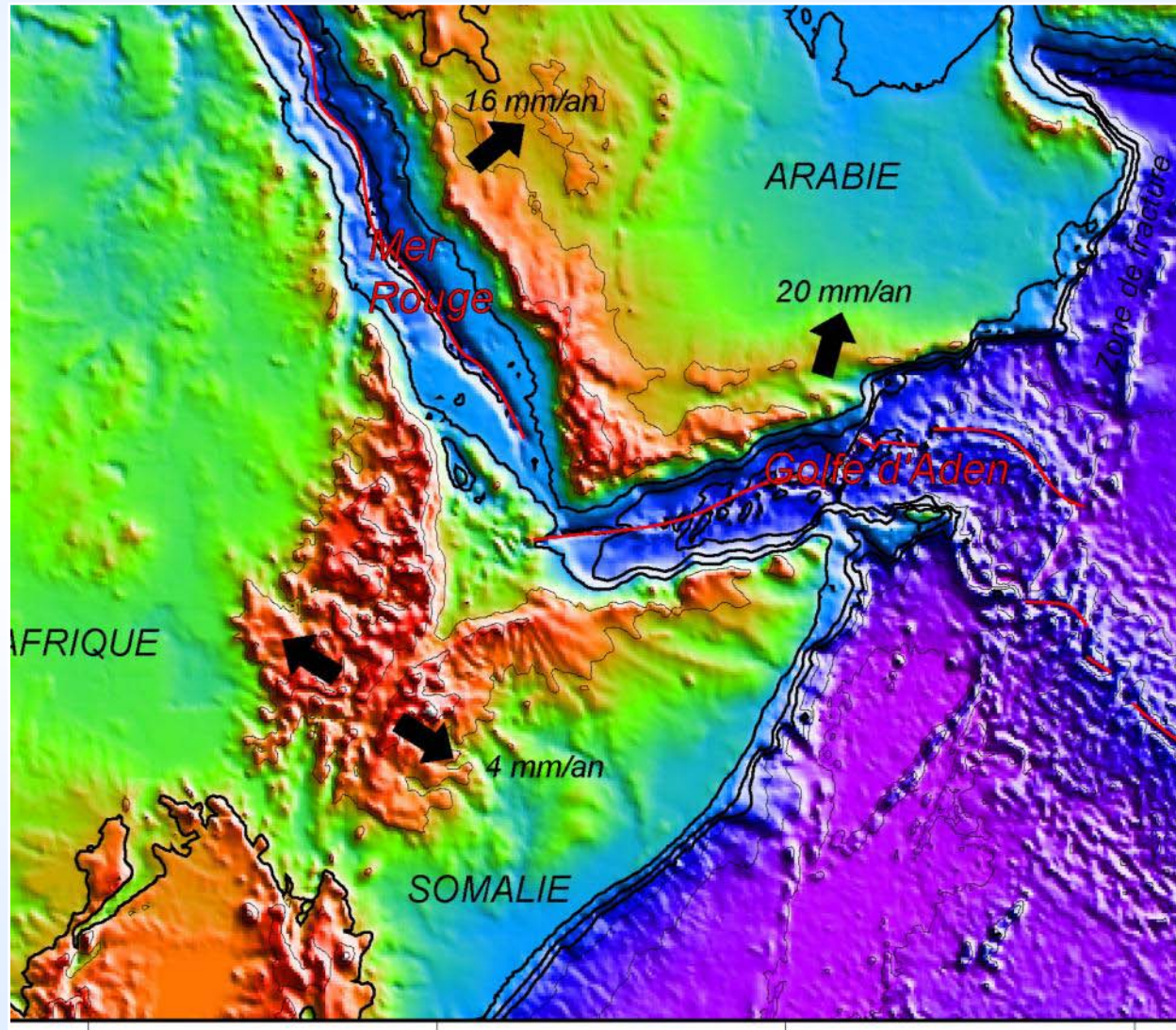
How do continents breakup?



Major conclusion

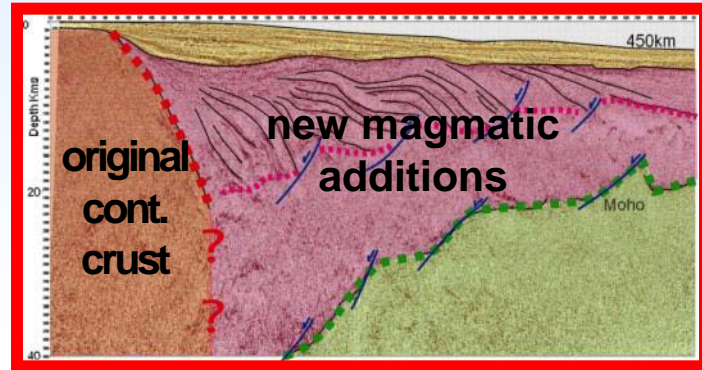
Most rifted margins may have undergone hyper-extension during earlier stages of rifting (also “magma-rich” systems), however, excess magma seems to be a prerequisite to breakup continental lithosphere (also in magma-poor margins)

How can we apply the lesson we learned in the Atlantic to the Afar system?

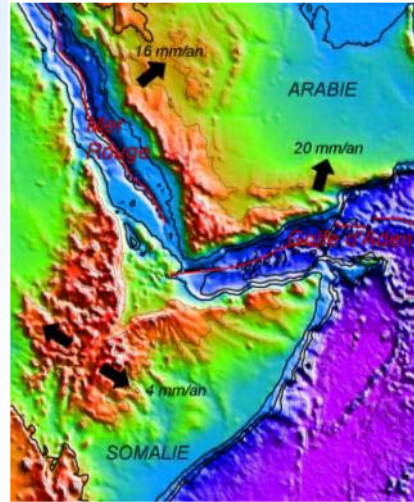


Two possible interpretations

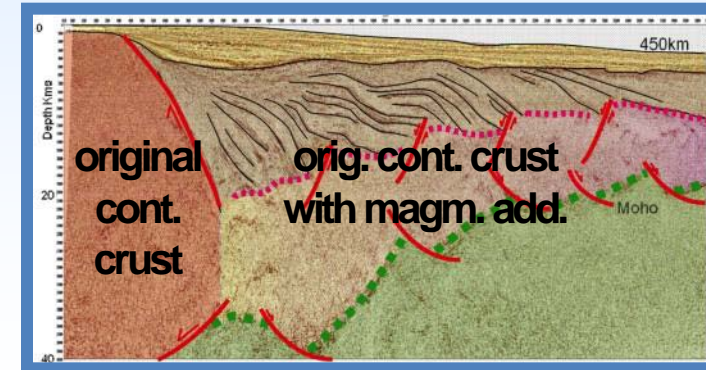
Purely magmatic



PelotasSPAN (ION-GXT) (taken from their homepage)



Hyper-extension with magmatic additions

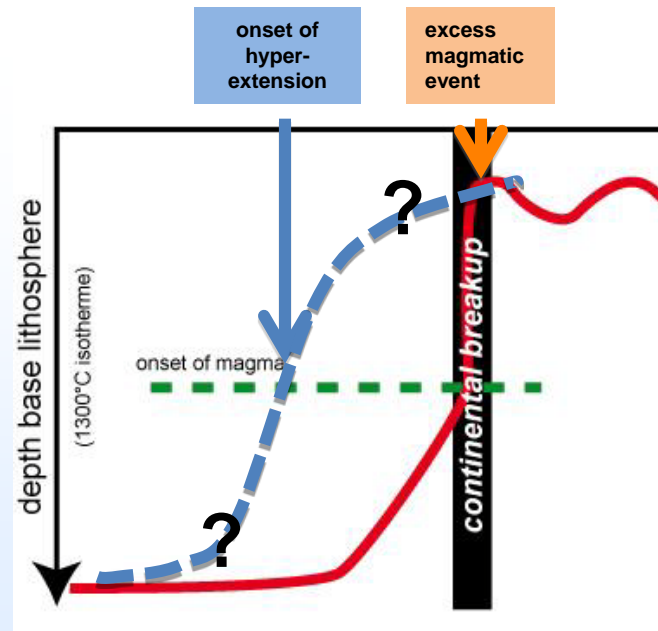


PelotasSPAN (ION-GXT) (taken from their homepage)

Interpretation I

- Whole crust is formed by newly formed magmatic crust (e.g. Island)

where are the margins and why the abrupt transition?



Interpretation II

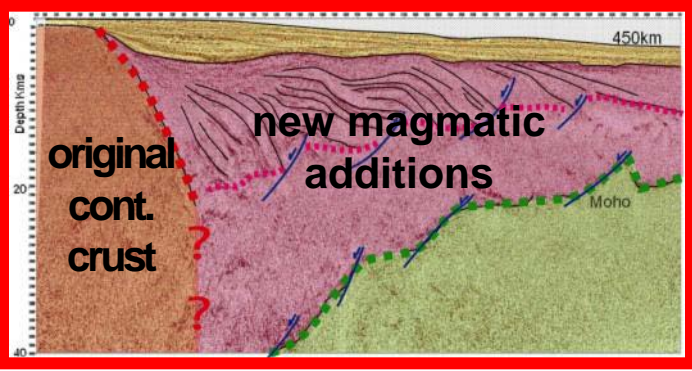
- Rifting initiates by hyper-extension
- Breakup is linked with excess magmatic event overprinting hyper-extended crust (e.g. San Paolo Plateau)

Is there evidence for hyper-extended crust and thinned lithosphere underneath AFAR ?

Two possible interpretations

Purely magmatic

Hyper-extensional magmatic



PelotasSPAN (ION-GXT) (taken from their homepage)

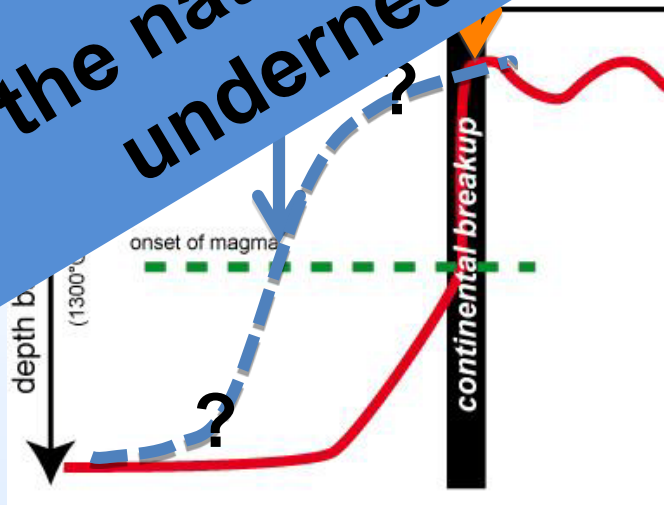
PelotasSPAN (ION-GXT) (taken from their homepage)

What is the nature of the crust/mantle underneath Afar?
 my questions:

Interpretation I

- Whole crust is formed by newly formed crust

(e.g. ...)
 where ...
 why the ...



Interpretation II

- Rifting initiates by hyper-extension
- Breakup is linked with excess magmatic event overprinting hyper-extended crust (e.g. San Paolo Plateau)

Is there evidence for hyper-extended crust and thinned lithosphere underneath AFAR ?