

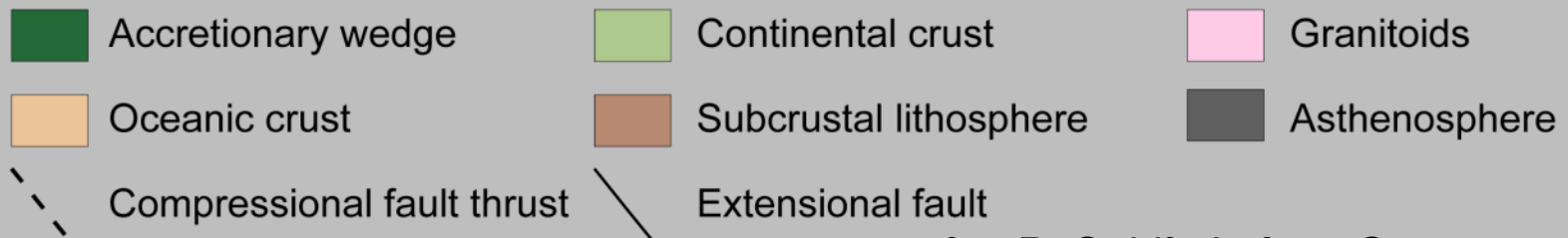
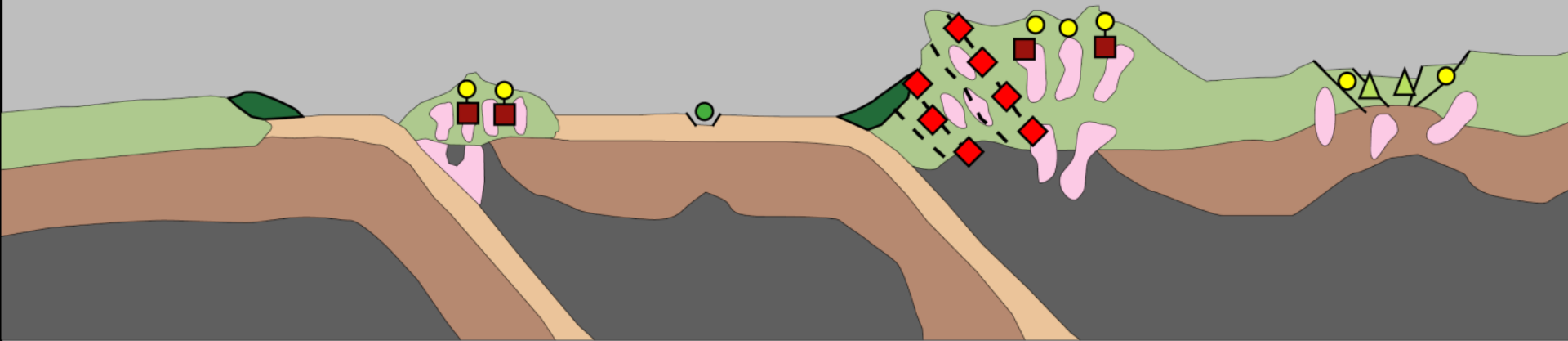
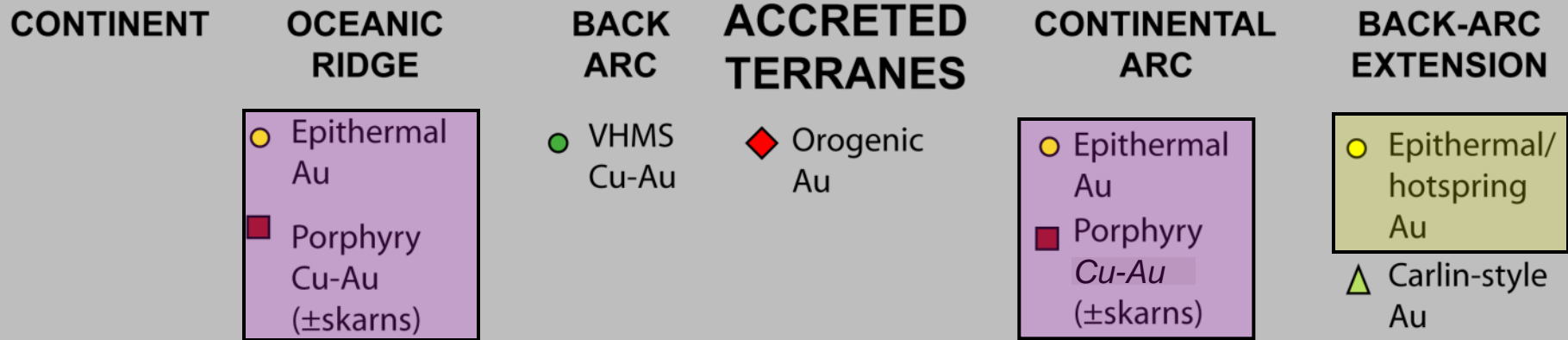
# Rift-related geothermal activity and epithermal gold veins

- NZ rift-related geothermal activity and characteristics
- Features of rift-related epithermal Au veins
- Controls on vein characteristics
- Potential for epithermal Au veins in Ethiopia

*with thanks to: Stratex East Africa Ltd.*

***Jeffrey W. Hedenquist***  
***University of Ottawa***

# Tectonic settings of hydrothermal systems



after R. Goldfarb, from Groves et al. (2005)

# ***Taupo Volcanic Zone, New Zealand***

***Looking NE: andesite arc to east,  
bimodal rhyolite-(basalt) in rift to west***

***~10,000 km<sup>3</sup>, ~1 M yrs:  
97.8% rhyolite, 2%  
andesite, 0.2% basalt***

***Mt. Tarawera rhyolite dome,  
late basaltic dikes***

***Mt. Tarawera:  
Pink and White  
Terraces  
(silica sinter)***



# Tarawera and Waimangu, NZ

*AD1315 rhyolite dome,  
driven by deep basalt*

Nairn et al., 2006

*Tarawera AD1886 dikes:  
basaltic fissure event*

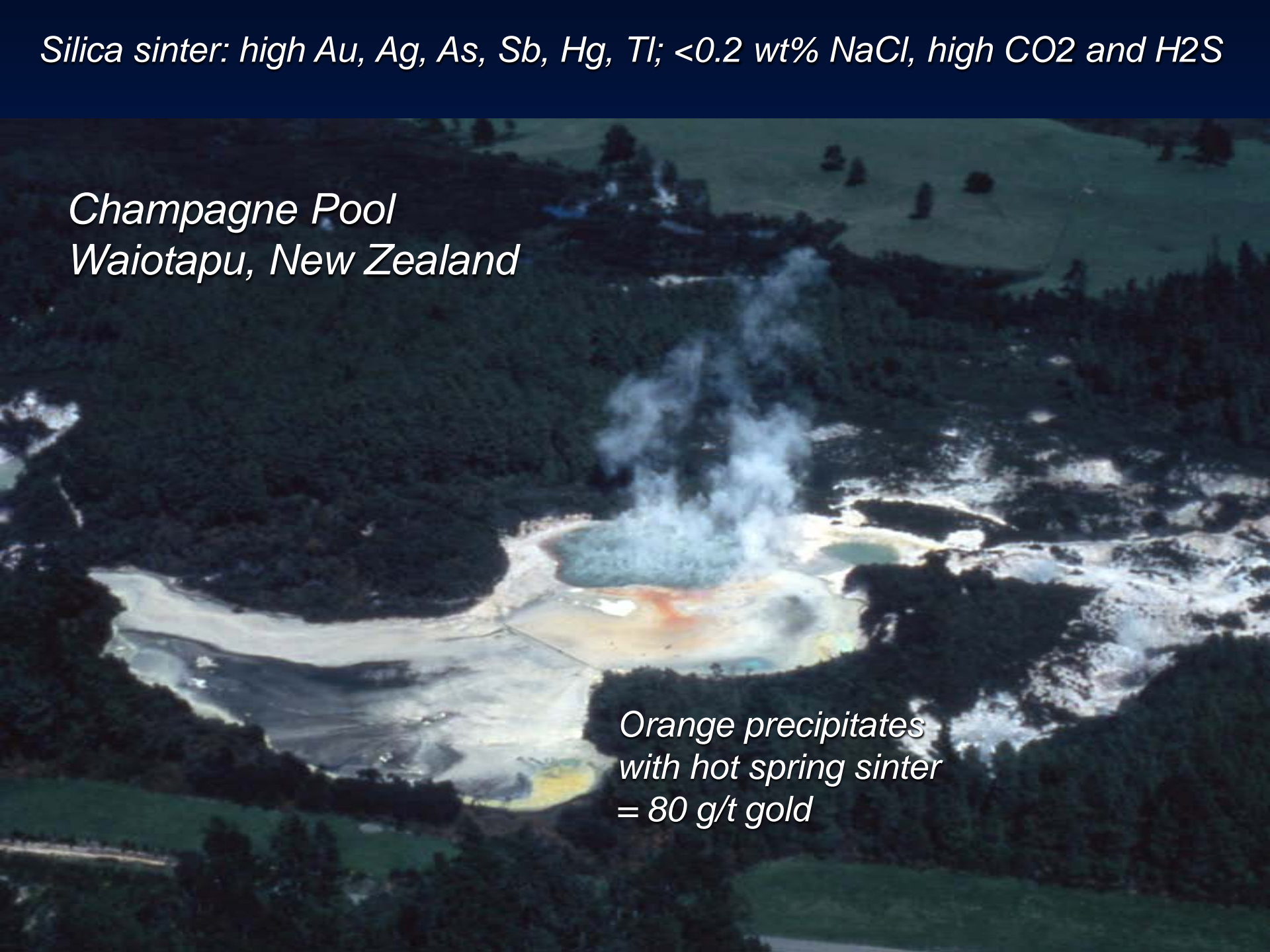
*Destruction of Pink  
and White Terraces*



*Silica sinter: high Au, Ag, As, Sb, Hg, Tl; <0.2 wt% NaCl, high CO<sub>2</sub> and H<sub>2</sub>S*

*Champagne Pool  
Waiotapu, New Zealand*

*Orange precipitates  
with hot spring sinter  
= 80 g/t gold*

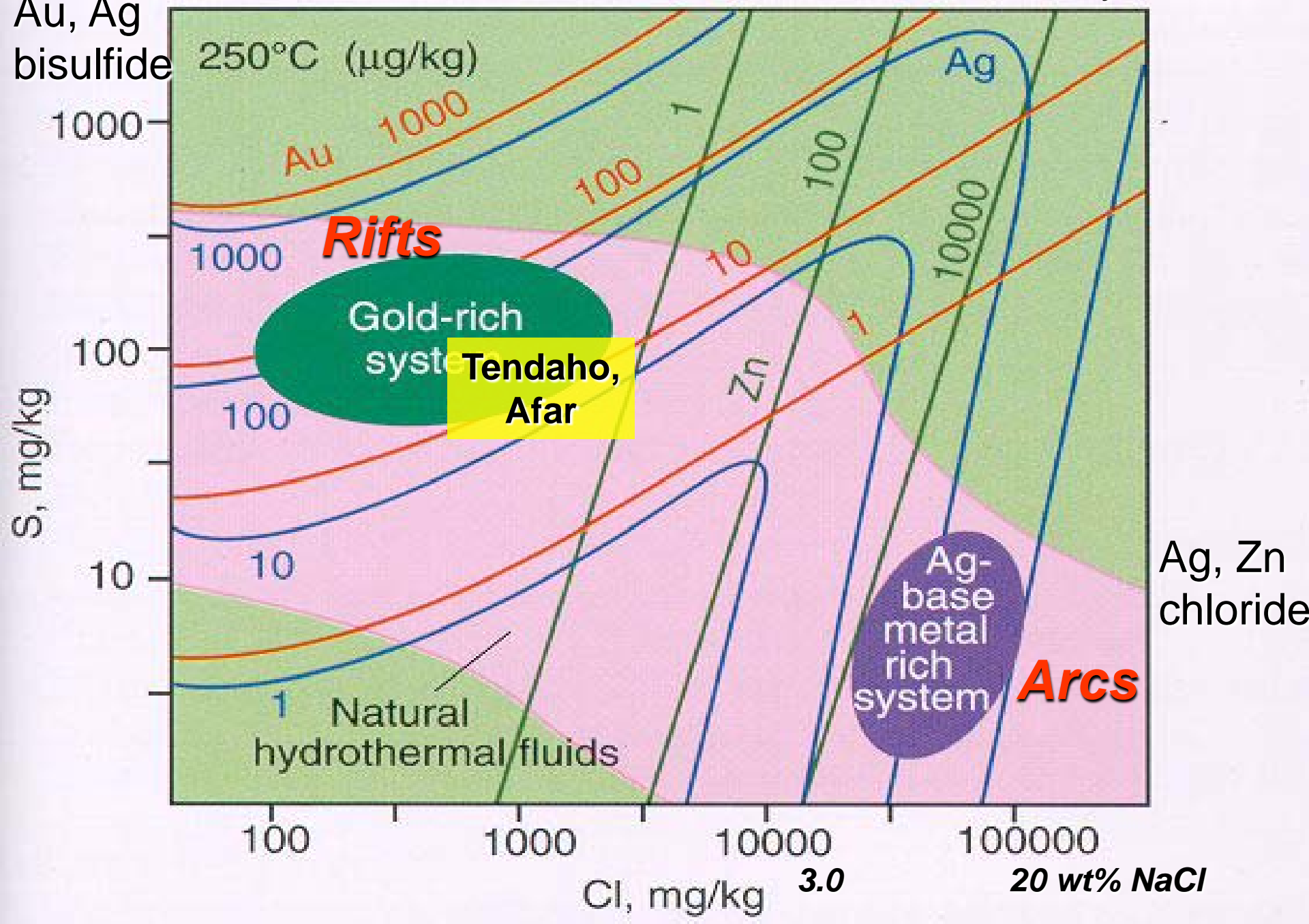
An aerial photograph of the Champagne Pool in Waiotapu, New Zealand. The pool is a large, irregularly shaped geothermal feature with a complex color palette. The central part of the pool is a vibrant orange, surrounded by yellow and green areas. The edges of the pool are fringed with white and grey sinter deposits. The surrounding landscape is a mix of dark green forest and lighter green grassy areas. The overall scene is a striking example of geothermal activity in a natural setting.

# Comparison of Taupo Volcanic Zone features with those of epithermal vein deposits

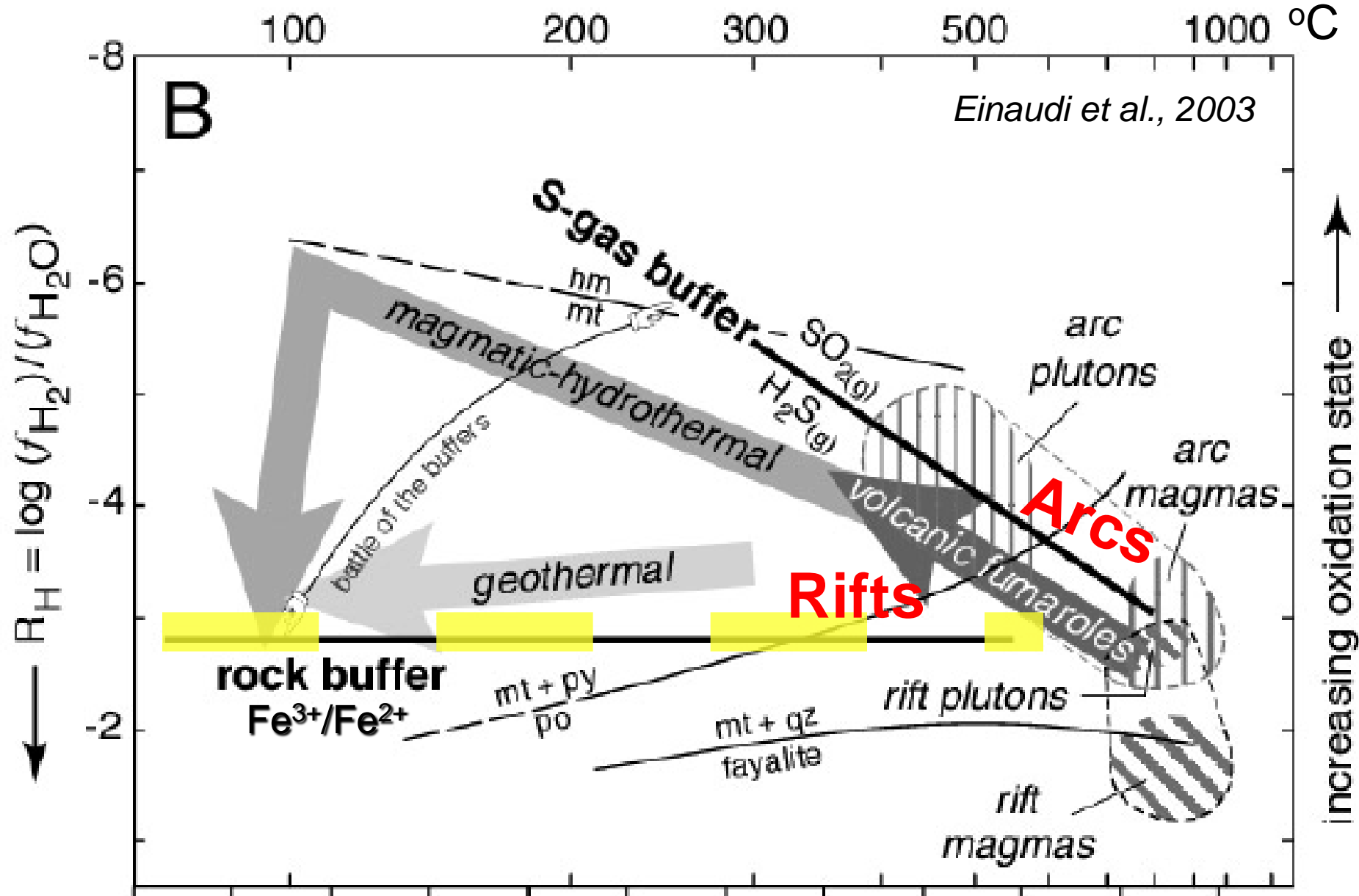
<b>Taupo Volcanic Zone</b>	<b>Back-arc extensional</b>	<b>Bimodal rhyolite-basalt</b>	<b>Au, Ag</b>	<b>&lt;0.2 wt% NaCl, lo to hi CO<sub>2</sub>, H<sub>2</sub>S</b>
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<b>Epithermal style</b>	<b>Tectonic setting</b>	<b>Magmatism</b>	<b>Metals</b>	<b>Fluids (NaCl, gases)</b>
Low sulfidation	Extensional; continental margin, back-arc or plume	Typically bimodal rhyolite-basalt, sub-alkaline	Au (Ag) Ag:Au ~ 1:1	0.1 to <2 wt% NaCl; CO <sub>2</sub> , H <sub>2</sub> S
Intermediate sulfidation	Neutral to extensional (or compressive) arcs	Andesite (dacite), calc-alkaline	Ag ± Pb, Zn (Au) Ag:Au > 20:1	3 to 7 (10 to 23) wt% NaCl; gases?

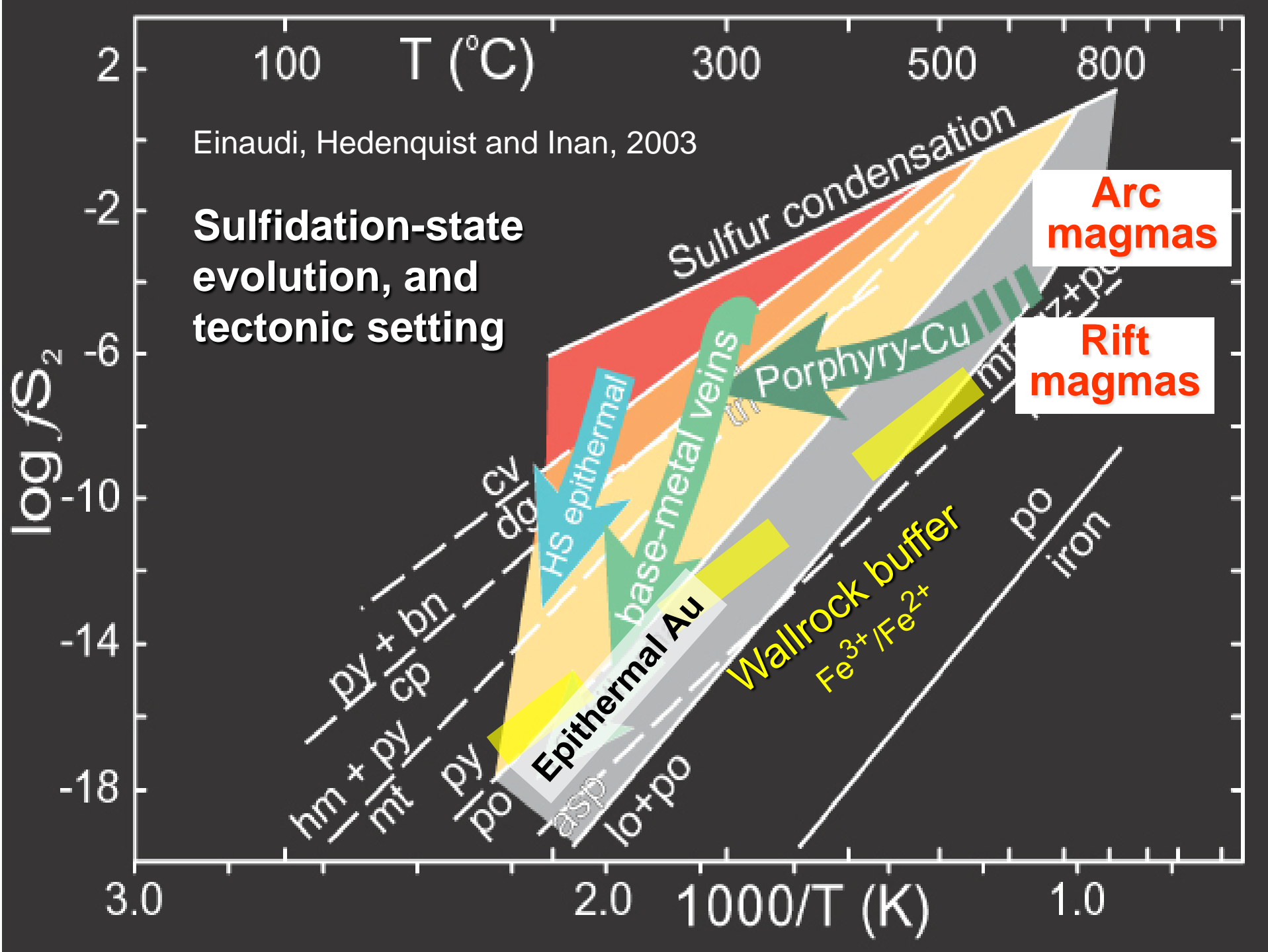
Au, Ag  
bisulfide



$R_H$  ( $\log f_{H_2}/f_{H_2O}$ ) vs.  $1/T$

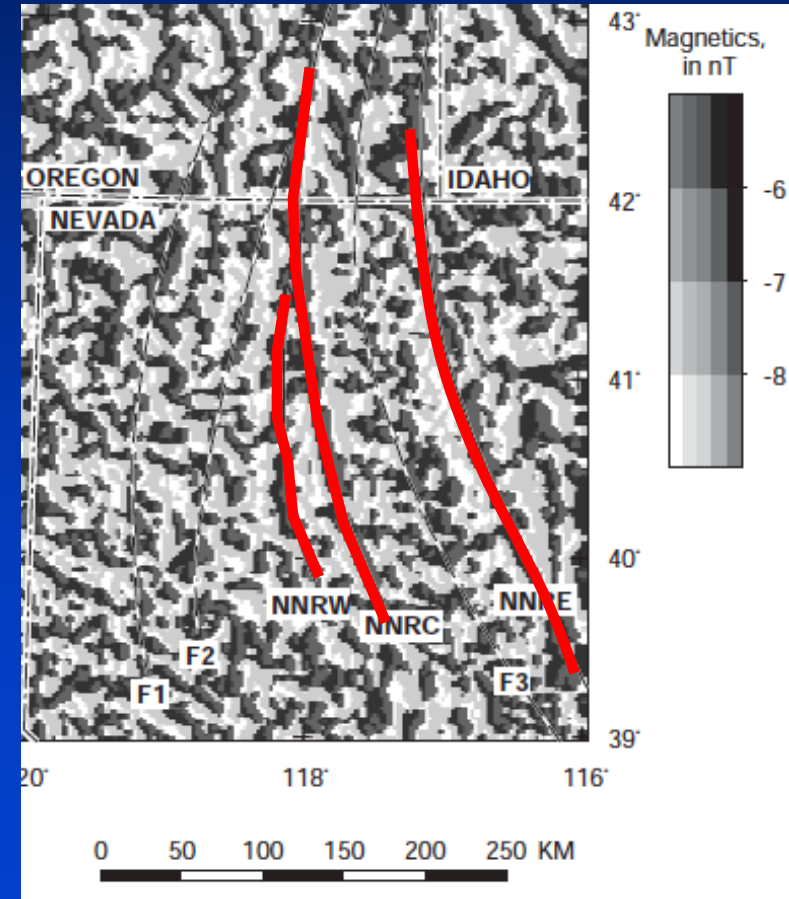
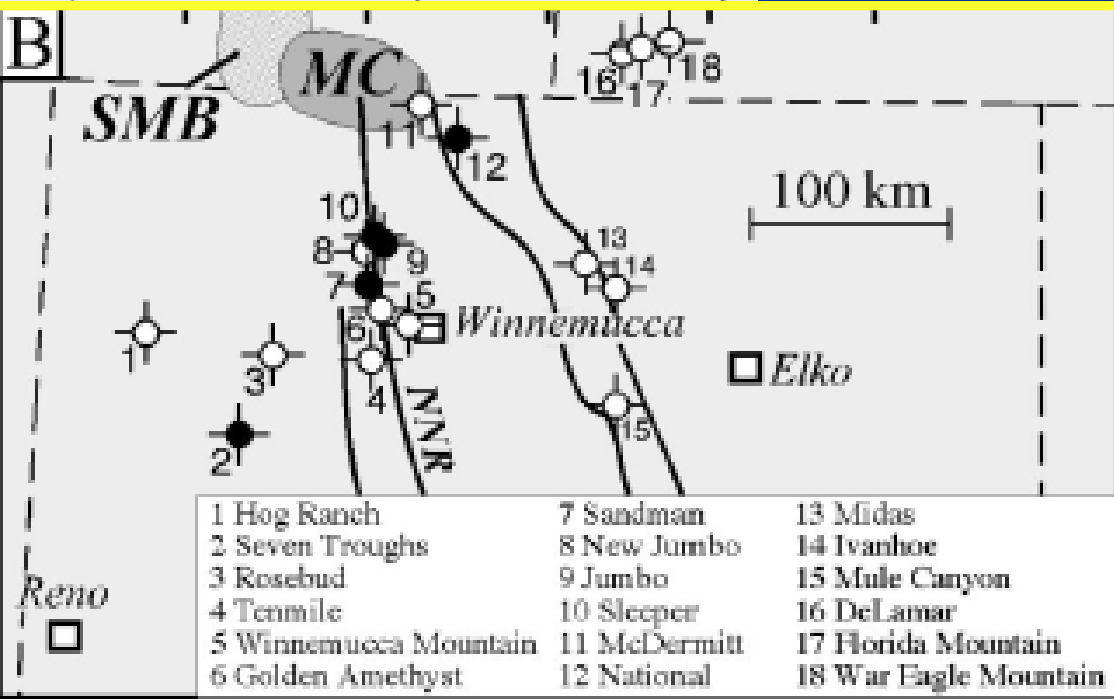
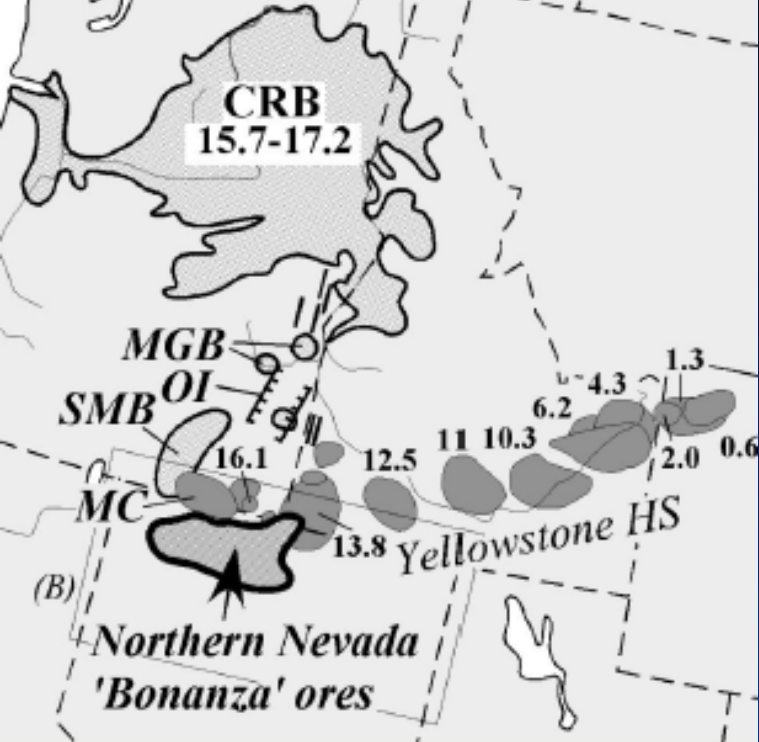


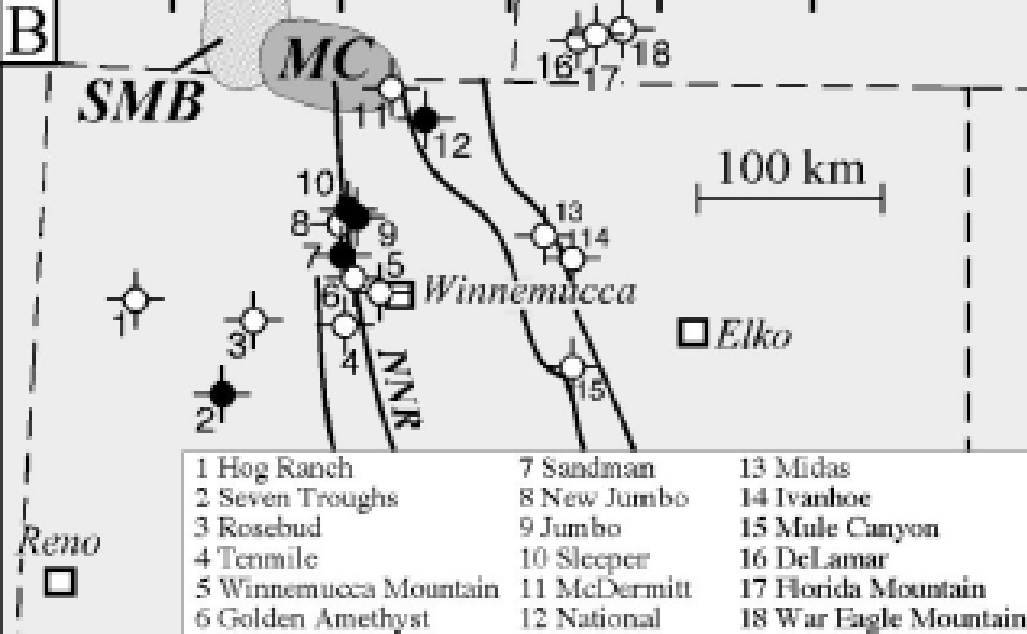




# Northern Nevada rift: Plume driven Rift-hosted epithermal Au deposits

*Geist and Richards, 1993; Dickinson, 2006  
Kamenov et al., 2007; Ponce and Glen, 2002*





# Northern Nevada rift and epithermal Au veins

*Saunders et al., 2010*

Sleeper low-sulfidation bonanza epithermal Au deposit, Nevada





Episodic events, extreme disequilibrium:

- Rapid fluid ascent, vapor loss*
- Formation of amorphous silica colloidal gels (colloform)*
- Vapor + gas loss (H<sub>2</sub>S, CO<sub>2</sub>)*
- Gold dendrite growth from colloids*



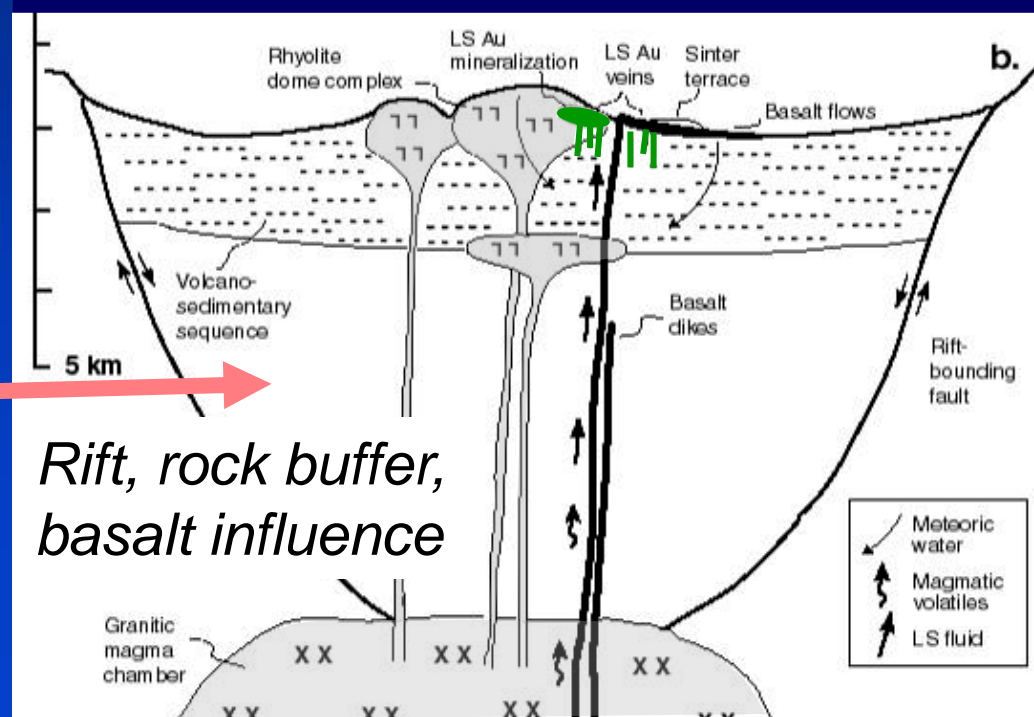
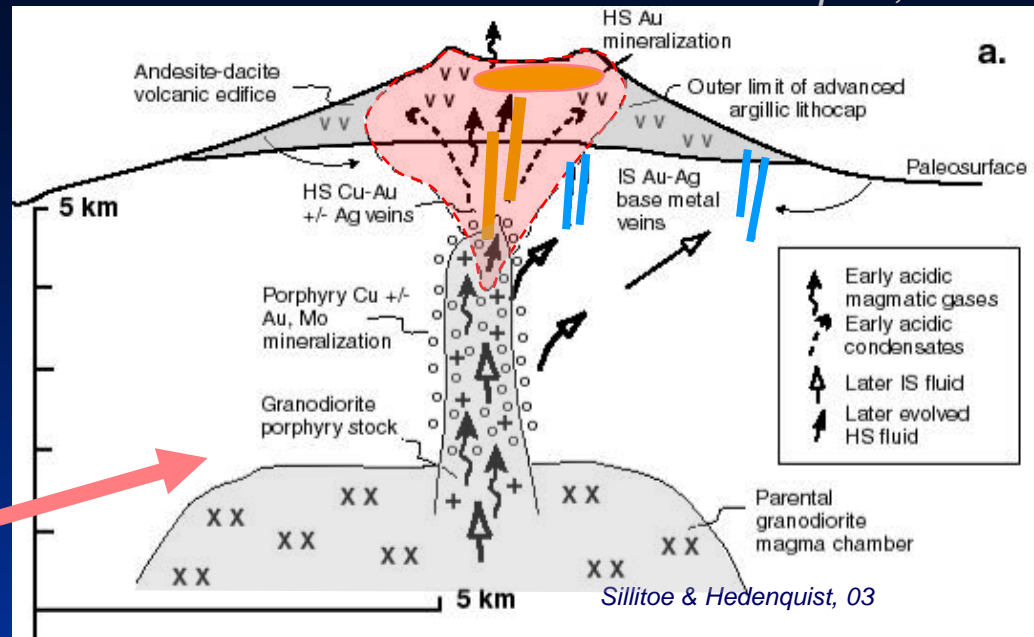
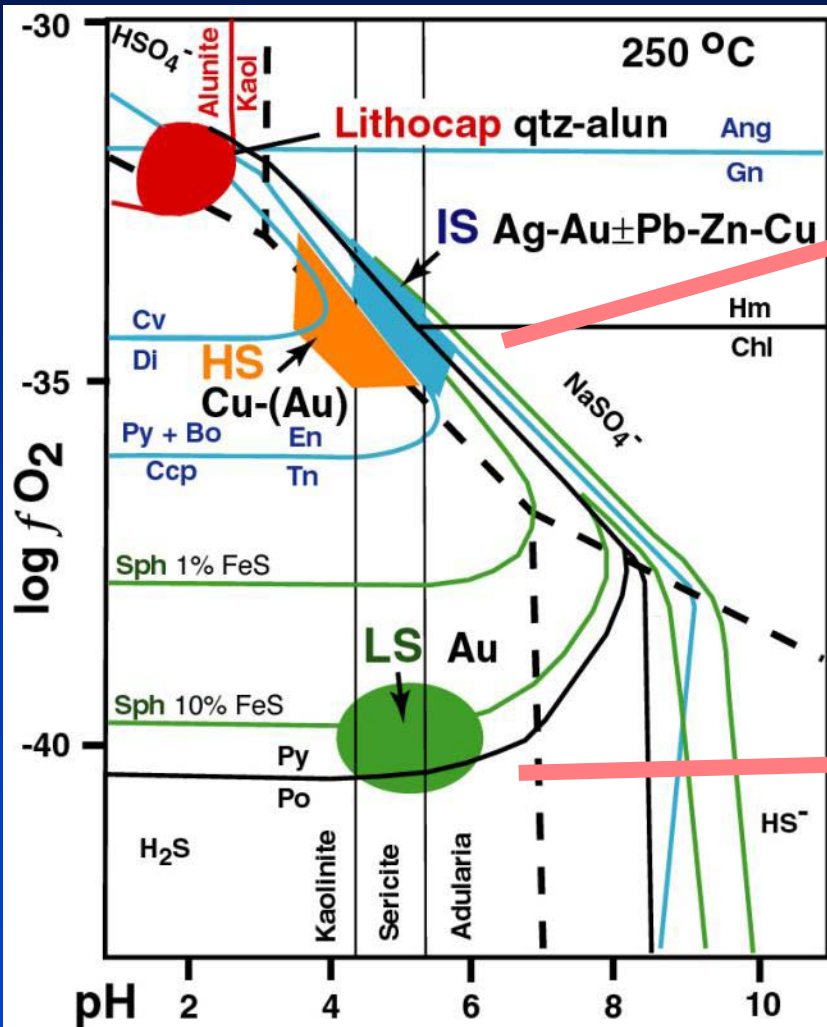
Colloform-banded quartz

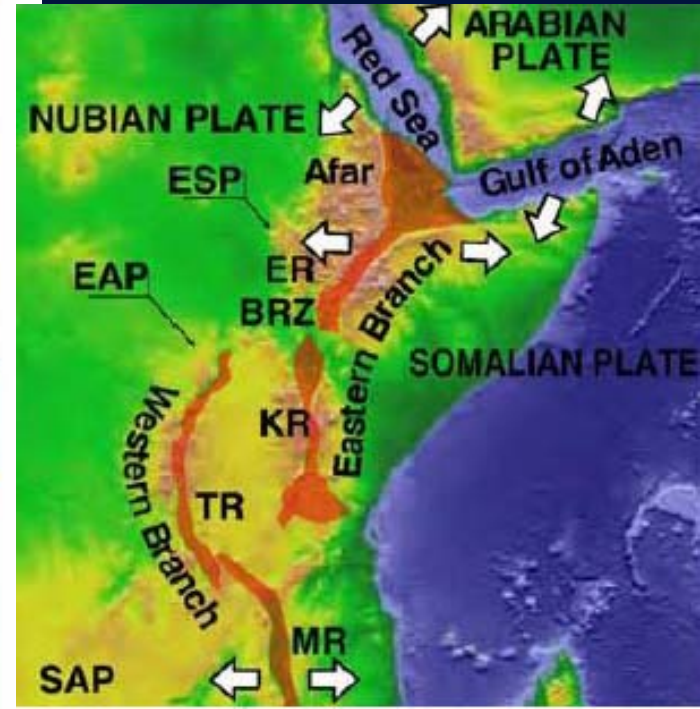
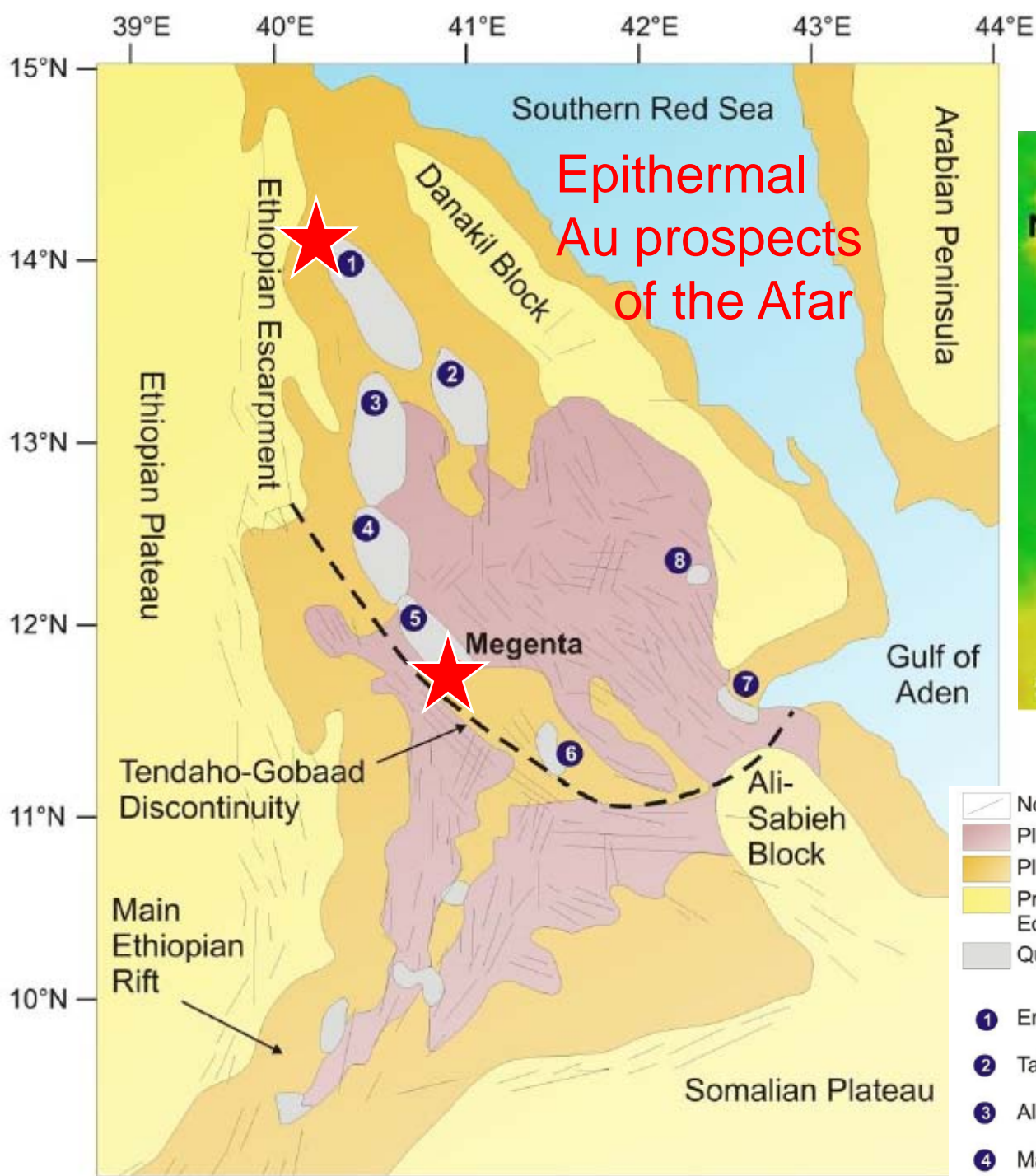


Dendritic Au, silica gels

*Sleeper*

# Epithermal deposit types: fundamental differences, arc vs. rift hosts



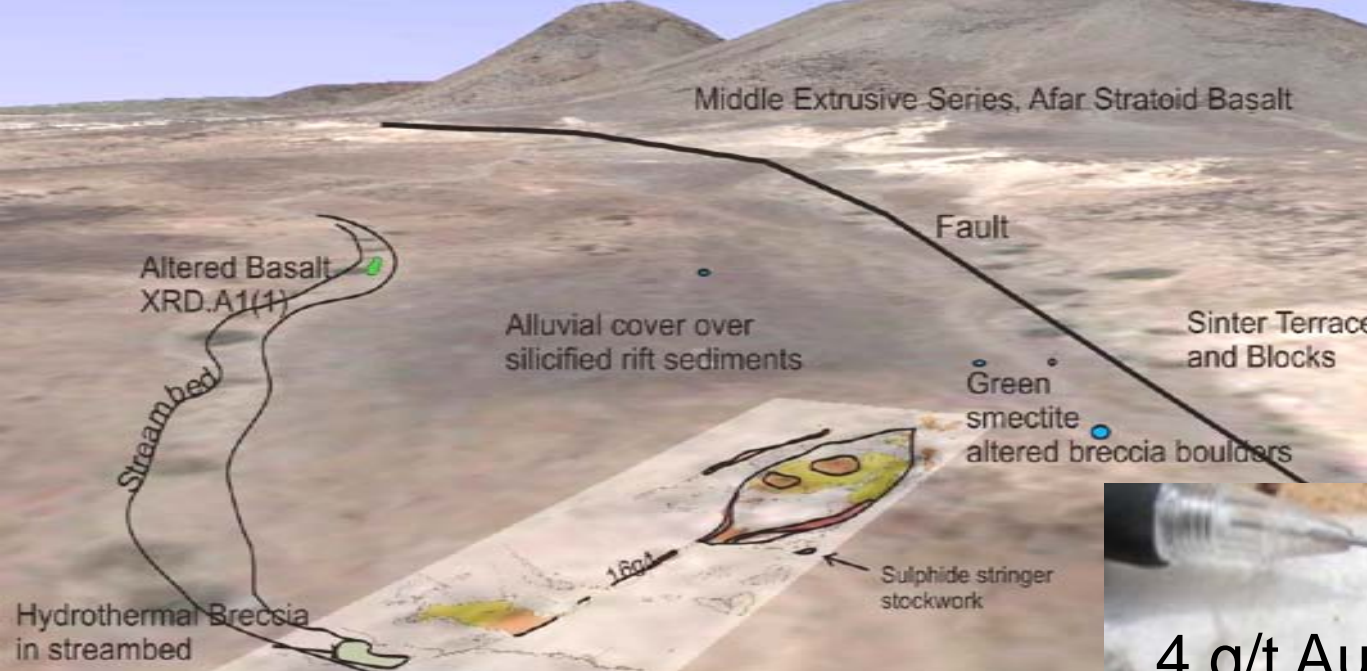


- Normal/strike-slip fault
- Pliocene flood basalts
- Pliocene - Recent sediments
- Proterozoic basement / Eocene-Miocene flood basalts
- Quaternary magmatic complexes

- |                                    |                             |
|------------------------------------|-----------------------------|
| <b>1</b> Erta'Ale volcanic complex | <b>5</b> Hararo rift zone   |
| <b>2</b> Tat'Ale volcanic complex  | <b>6</b> Gobaad rift        |
| <b>3</b> Alayta volcanic complex   | <b>7</b> Asal-Ghoubbet rift |
| <b>4</b> Manda rift zone           | <b>8</b> Manda-Inakir rift  |

# Magenta epithermal Au prospect, Tendaho

Lavelle, MSc 2010;  
Stratex, 2011

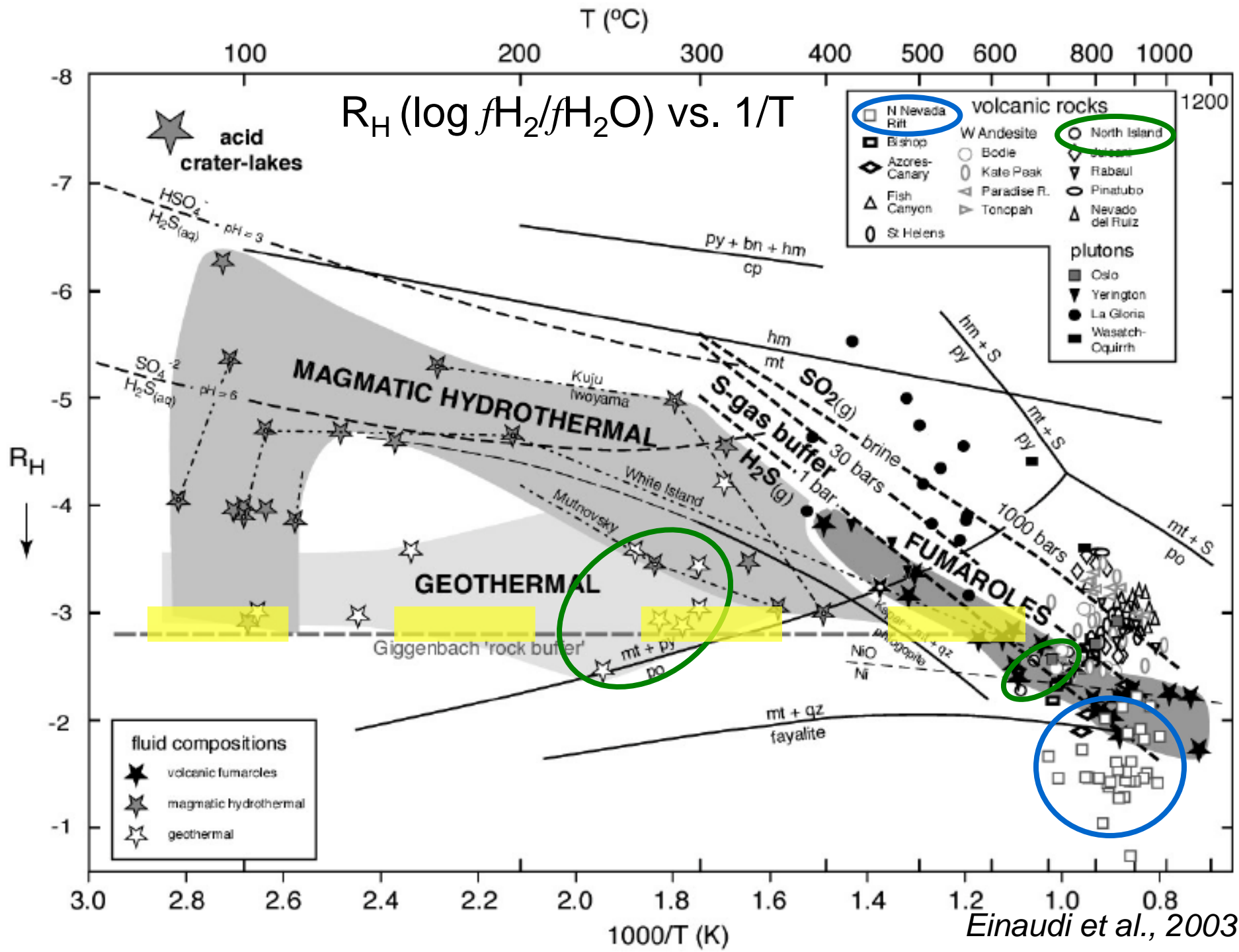


# Conclusions: Volcanic-hosted epithermal Au veins

- **Rift-related geothermal systems:**
  - Thick rhyolite sequences; deep basaltic magma
  - Relatively reduced magmas and fluids; rock buffer control
  - Low salinity but variable (high) gases (particularly H<sub>2</sub>S)
  - Strong fault-related permeability, episodic opening
- **Low-sulfidation epithermal Au veins:**
  - Rifts: rhyolite-basalt bimodal products, dome hosts
  - Rock buffer control, low salinity; H<sub>2</sub>S rich: high Au
  - Rapid ascent, boiling: Au and silica supersaturation
  - Basins, low relief: silica sinter aprons (steam-heated clays)

**Ethiopia: right setting for bonanza Au veins**





$R_H$  ( $\log [fH_2/fH_2O]$ ) vs  $1/T$

