

Controls on the Frequency of Dike Episodes At Spreading Centers



Buck, January, 2006

Dabbahu, Ethiopia
a



Thorarinsson September 8, 1977

Krafla, Iceland

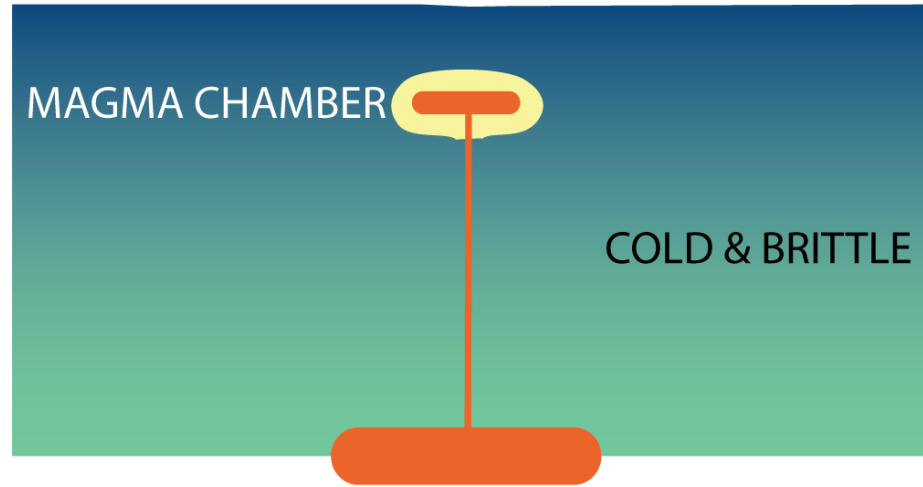


Fonari and Shanks, May, 2006

EPR, 9° N

'Bottom Up'

driving of dikes requires magma overpressure

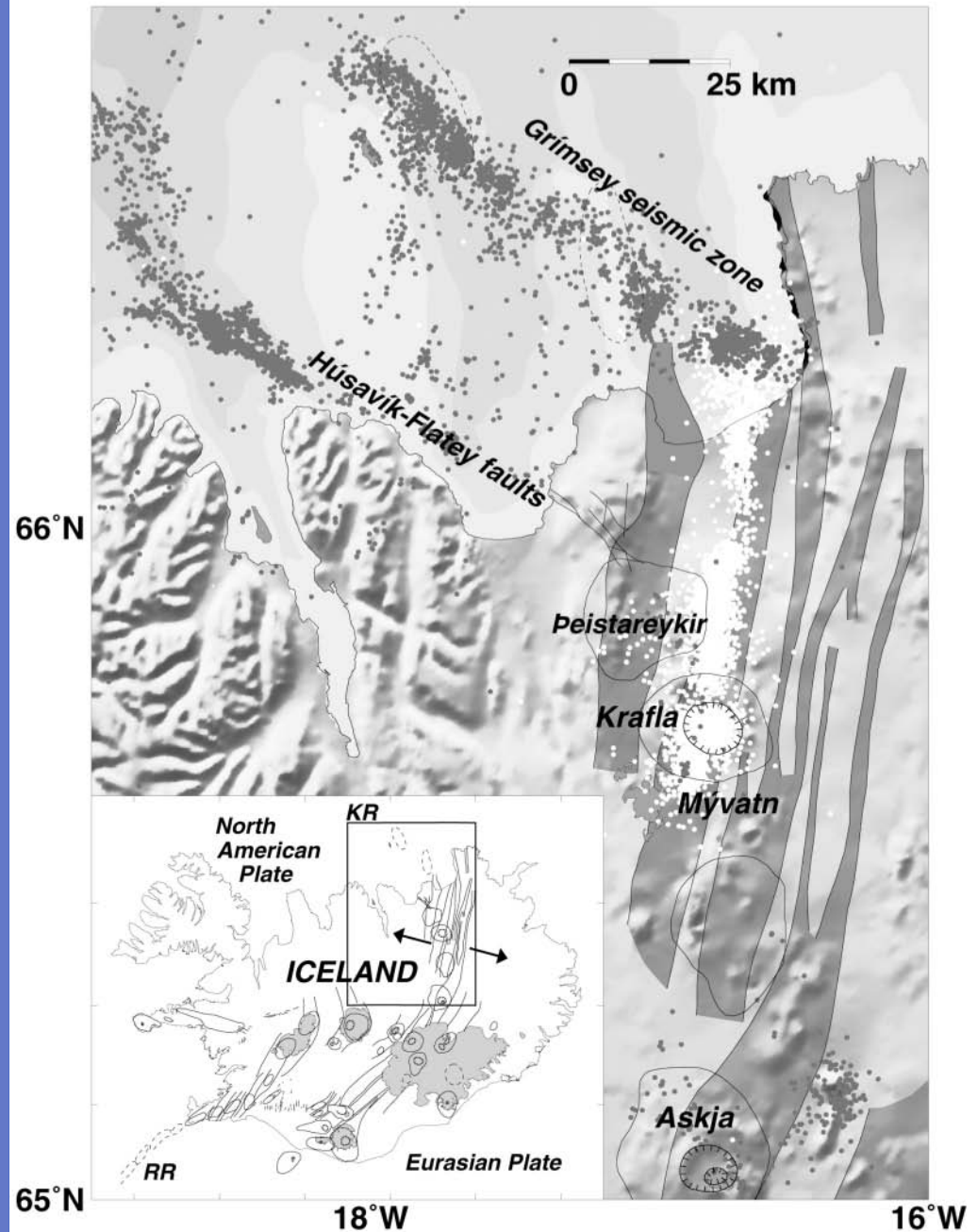


'Top Down'

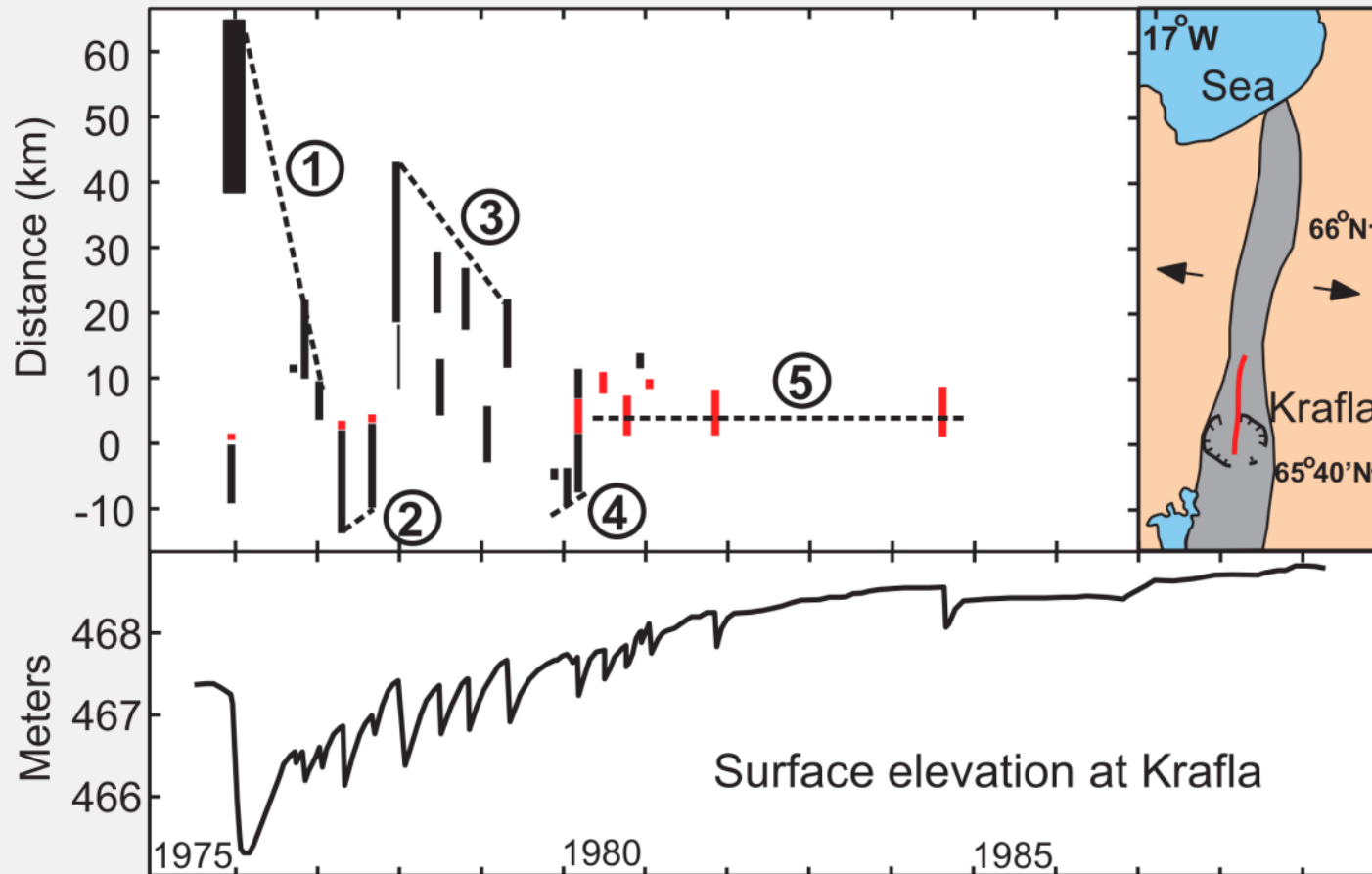
driving of dikes tectonic stress



Krafla, Iceland
Earthquake and Fissure Locations
(by Bryndis Brandsdottir)

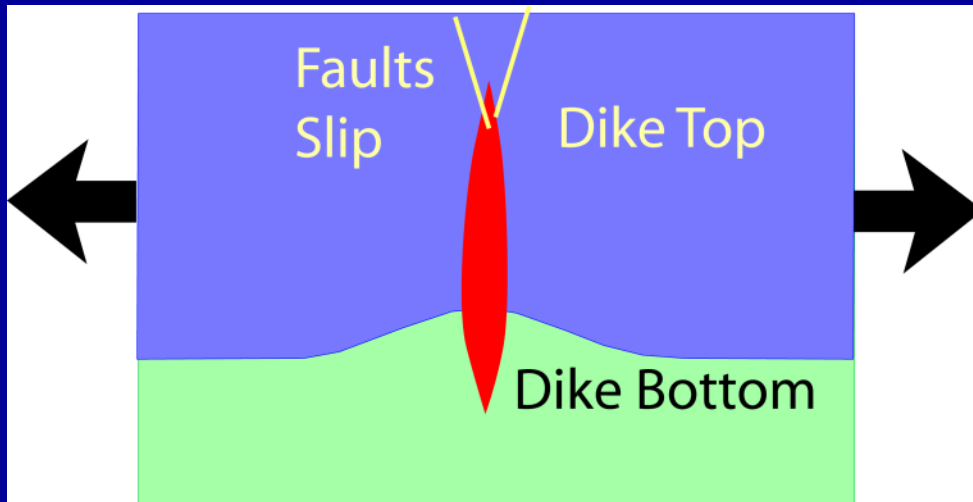


1975-1984 Krafla Rifting Episode

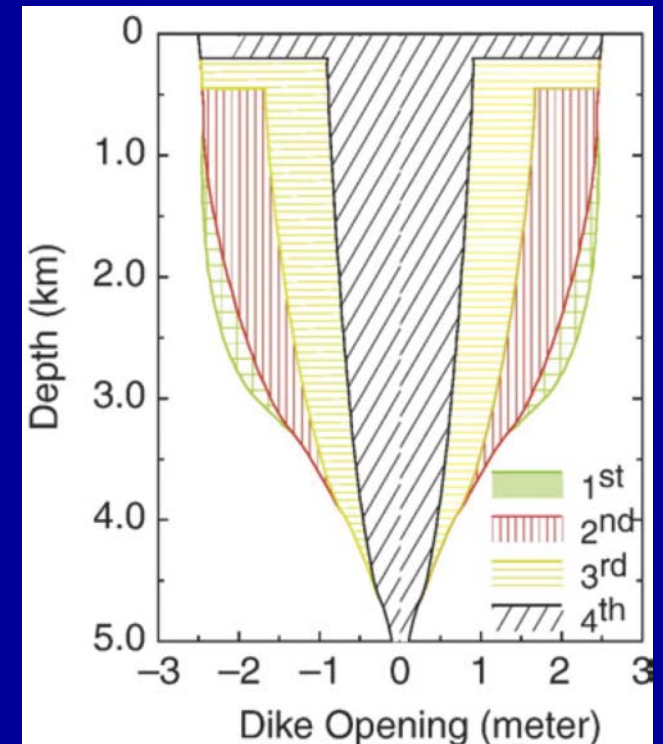


(Adapted from Einarsson, 1991)

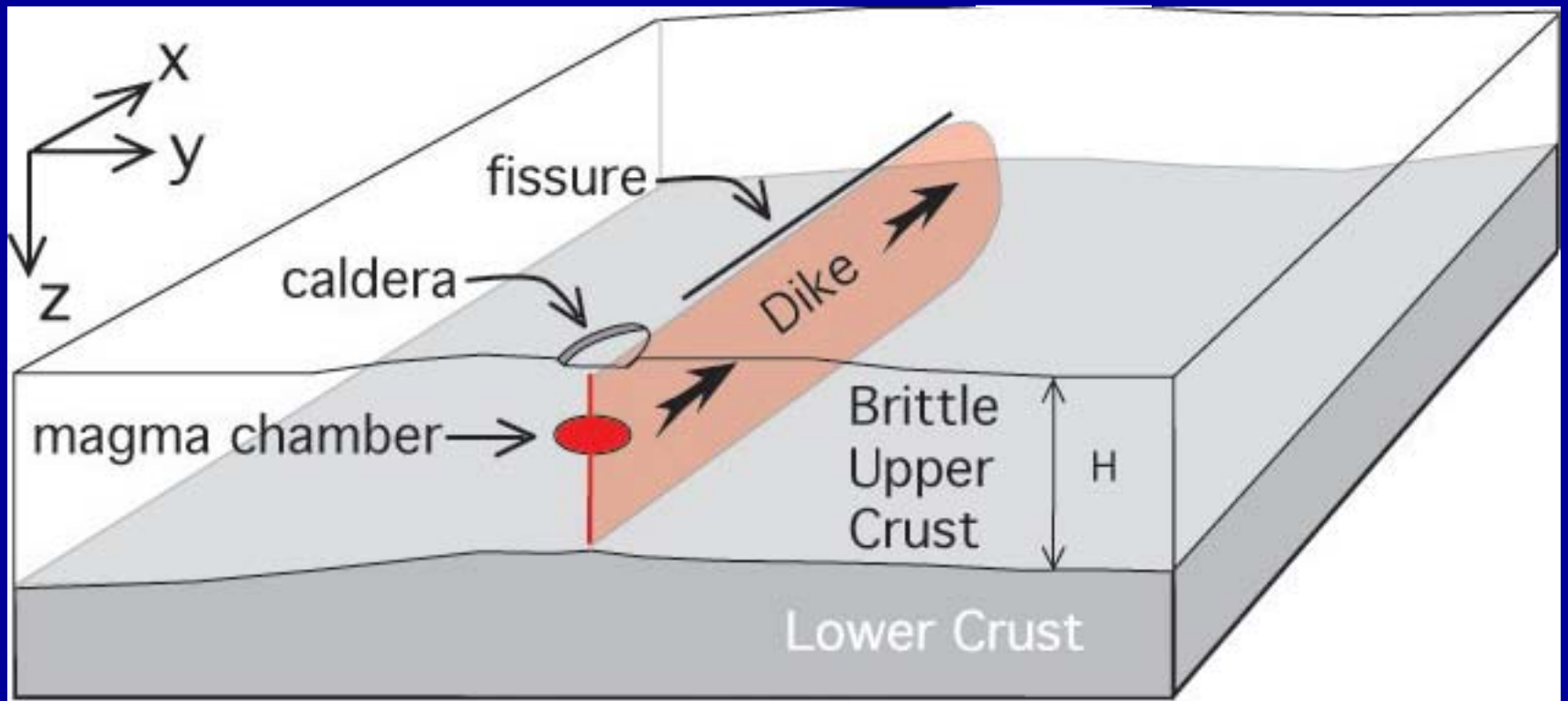
Stress Release by one dike makes next dike shallower



After Rubin and Pollard (1987)



Qin and Buck, 2007



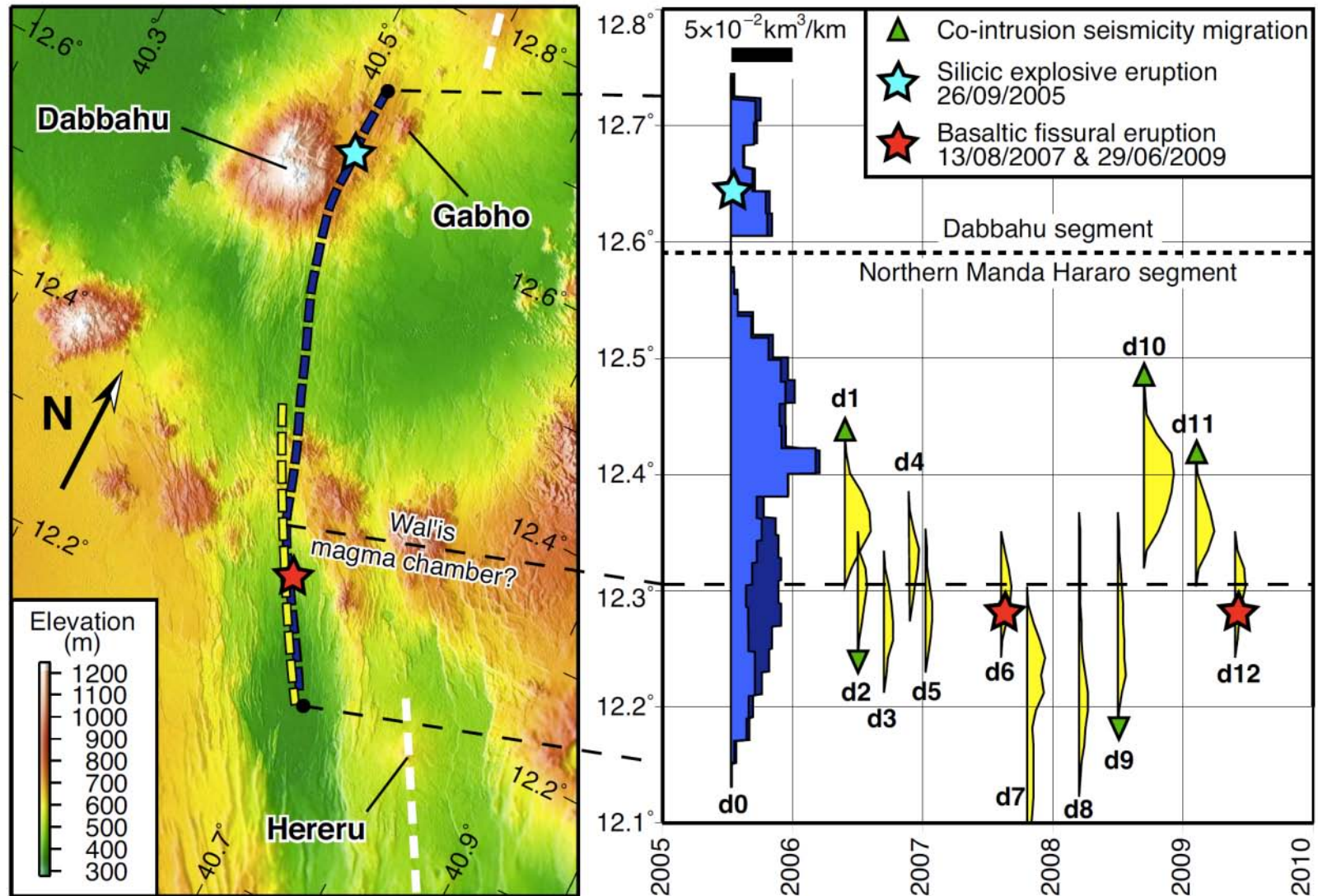
Inferences from Krafla

1. Micro-seismicity before first event
2. First event was different from other events
(e.g. wider, hit segment end)
3. Events continue until tectonic stress relieved
4. Time interval between recent episodes 250 years,
so extension since last event was 5 m.
5. Depth of magma chamber is 3 km.

Dabbahu



(2010)

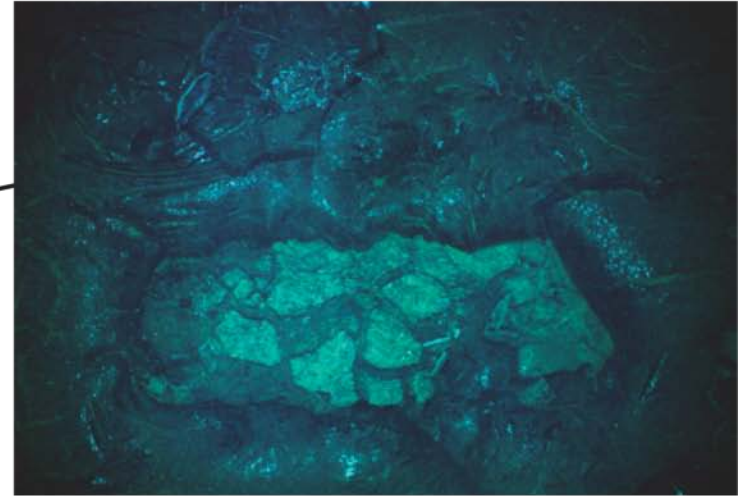
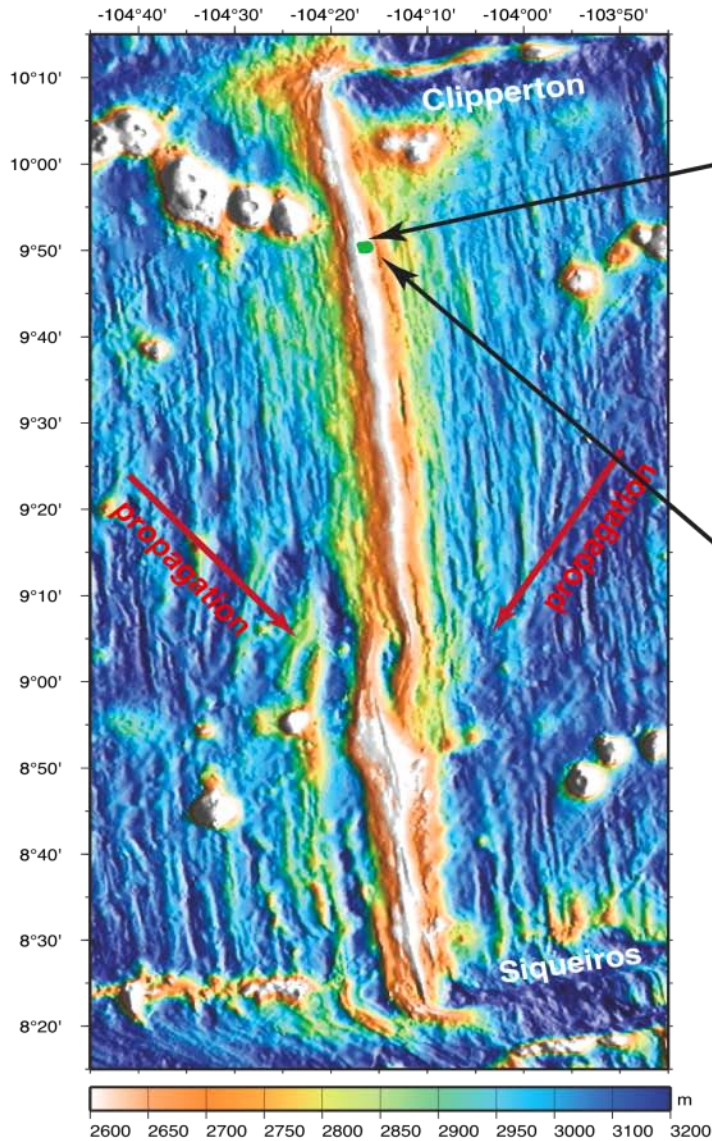


Inferences from Dabbahu

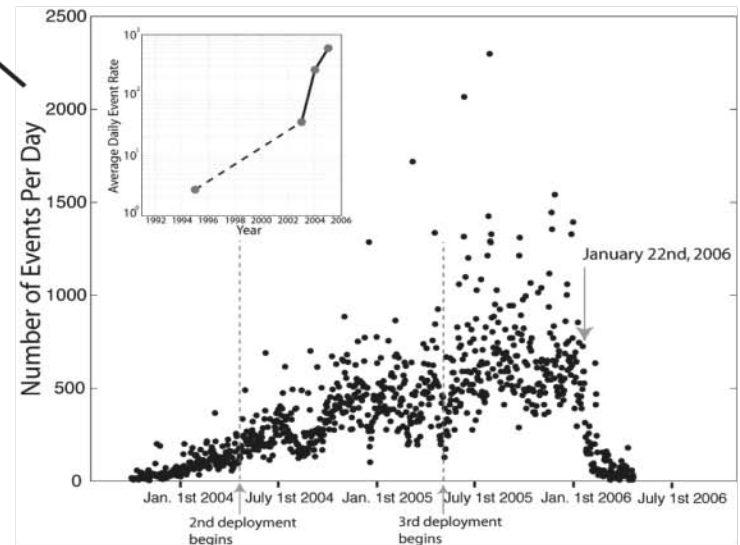
1. Micro-seismicity not known
2. First event was different from other events
(e.g. much wider, went farther)
3. Events continue until tectonic stress relieved?
4. Time interval between recent faulting episodes
1700 +/-?? Years, opening >10m.
5. Depth of magma chamber is unknown.

9°N East Pacific Rise

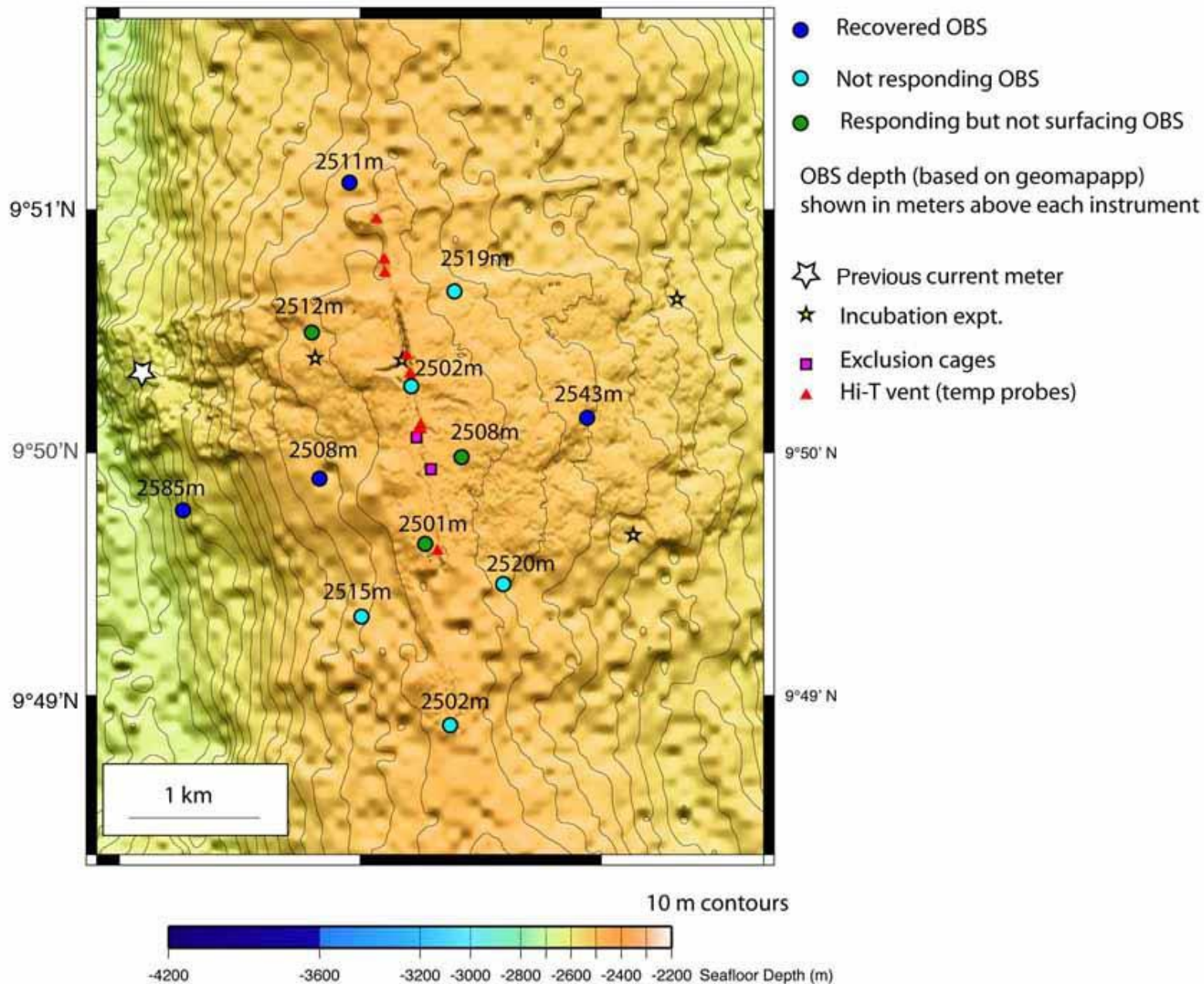
~ 15 Years Between Fissure Eruptions
at 9° 50' on the East Pacific Rise



Fresh Glassy Basalt at 9° 47'N taken May 2006 by Fornari and Shanks



2003-2006 microseismic events from Tolstoy et al. (2006).
1994 average event rate point from Sohn et al. (1999).

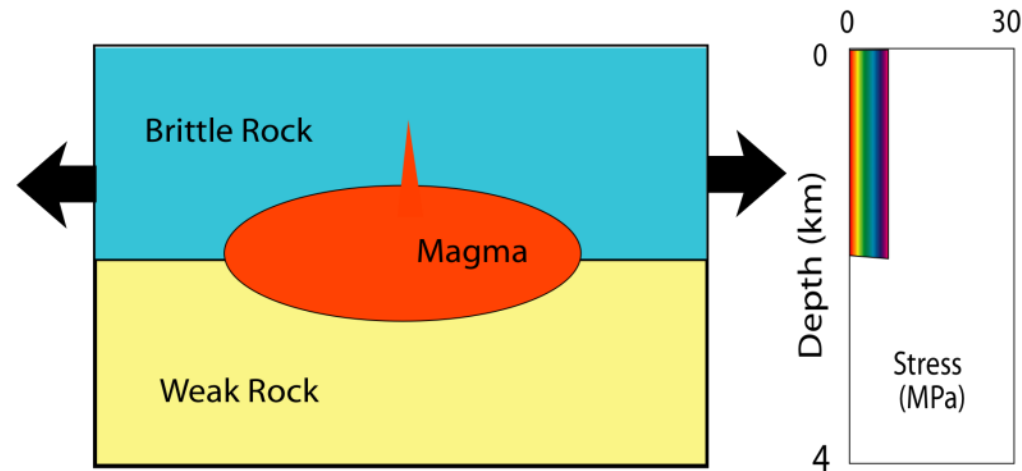


Inferences from 9° 50'N EPR

1. Rising micro-seismicity before first event
2. No information on number of events
3. Lava flows indicate that tectonic stress relieved
4. Time interval between recent episodes 15 years, so extension since last event was 1.9 m.
5. Depth of magma chamber is ~1 km.

Old Model

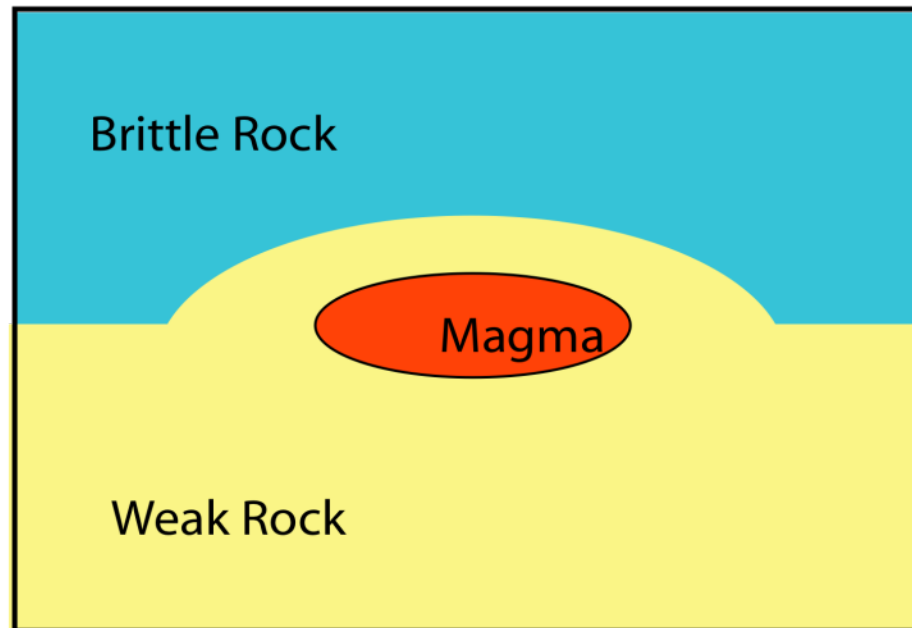
Noting that most dikes formed at spreading centers are about a meter wide a simple model for dike frequency has been widely discussed (e.g. Gudmundsson, 1993). Spreading center magma chambers are assumed to be in contact with the brittle upper crustal lithosphere that is stressed by plate motion.



Dike intrusion requires low viscosity magma in contact with elastically stressed rock. Stress builds so that a 1 meter-wide dike can form.

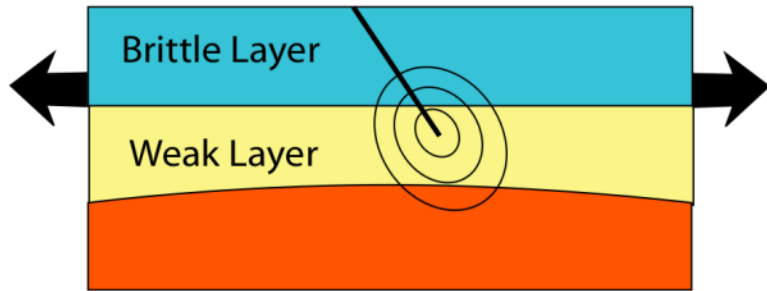
New Model

For quasi-steady-state magma chambers at spreading centers there may be a "buffer zone" of weak rock between fluid magma and brittle rock

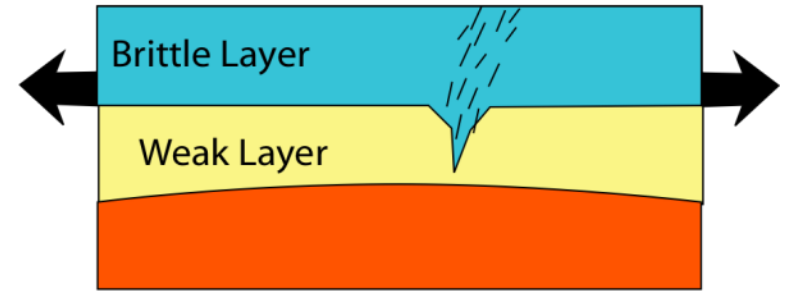


To get dikes may require earthquakes to connect magma and stressed rock.
This may occur in different ways.

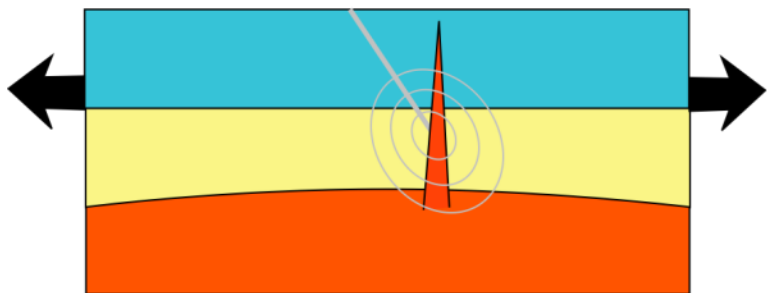
A. Big Earthquake Stresses Buffer Zone



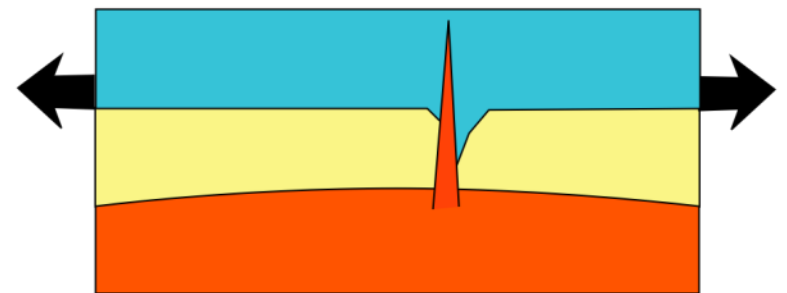
B. Many Small Earthquakes Leads to Hydrothermal Thinning of Buffer Zone



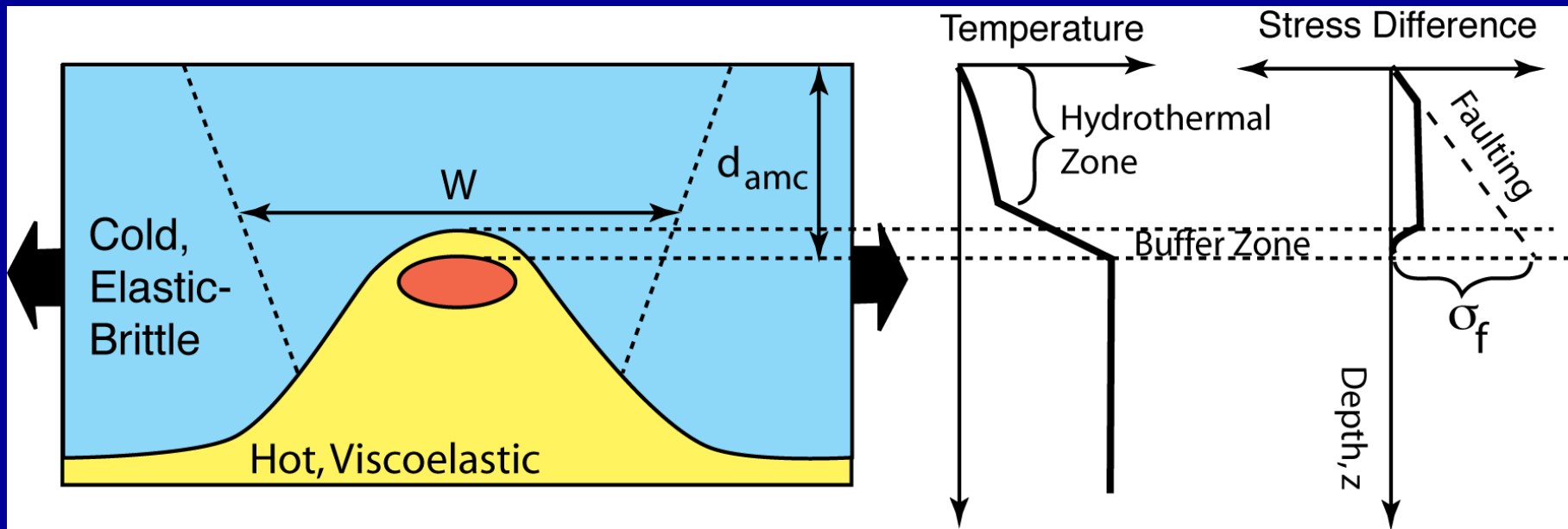
Dike Triggered as Weak Buffer Cracks

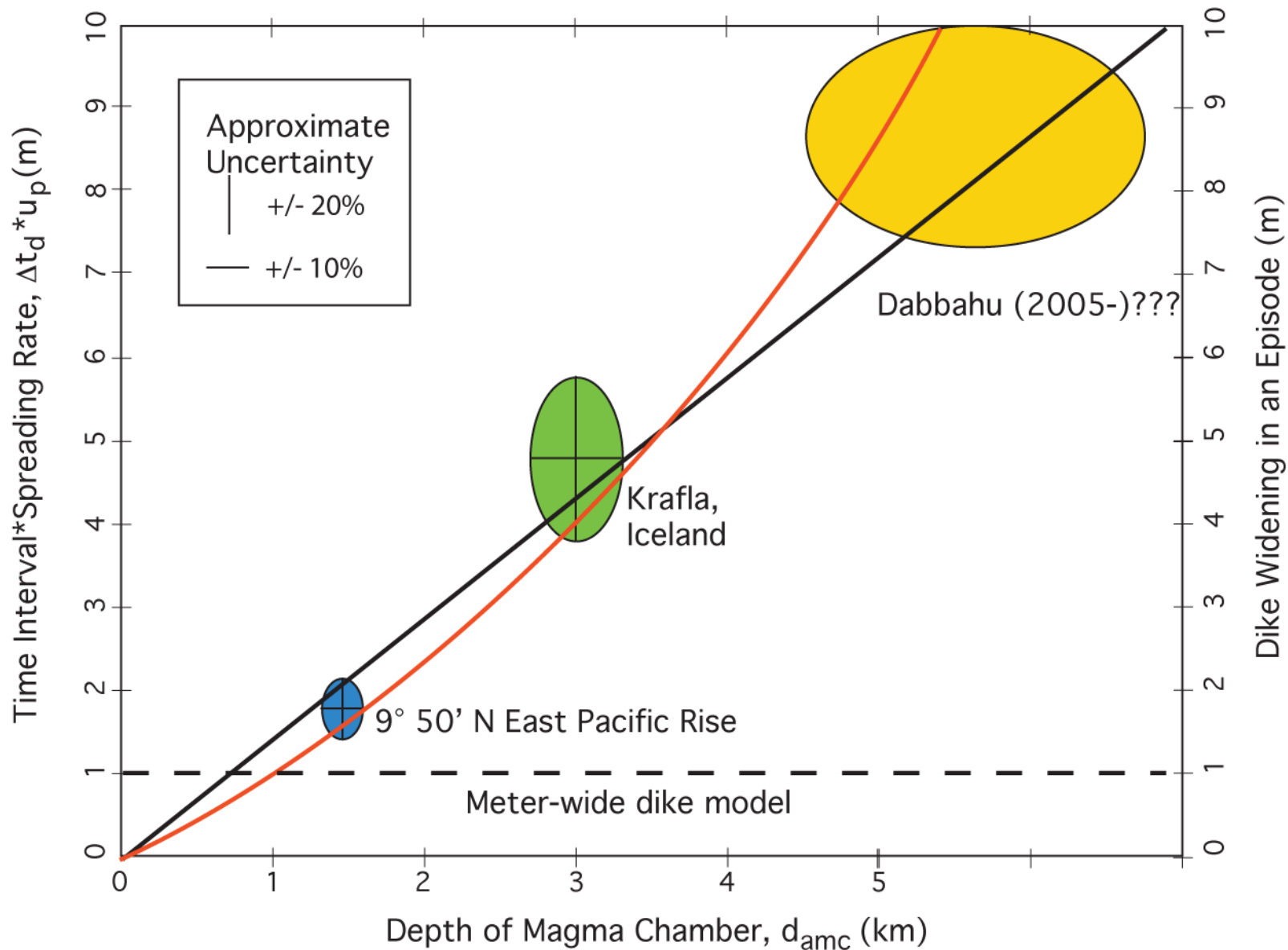


Dike Triggered as Small Quake Cracks Buffer



If
Faulting is needed to trigger diking
and
A dike episode relieves stress differences
then
We can estimate the time between Dike Episodes





Conclusion

Fault triggering of spreading center dikes explains:

1. Frequency of dikes at different spreading centers
2. Onset and buildup of seismicity before diking